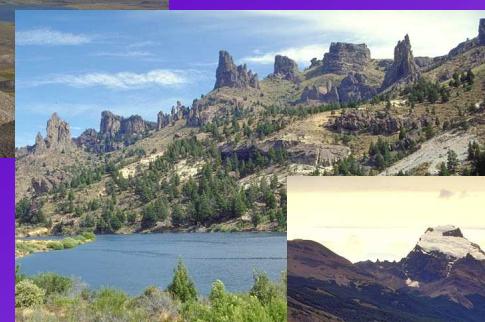
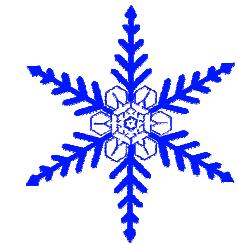


Anomalías de la Circulación Atmosférica asociadas a extremos de caudal en ríos andinos y posibilidades de pronóstico a mediano plazo



Dr. Diego C. Araneo

IANIGLA – CONICET
CCT - Mendoza

Características de la variabilidad espacio-temporal de los caudales de los ríos de la región

OBJETIVO:

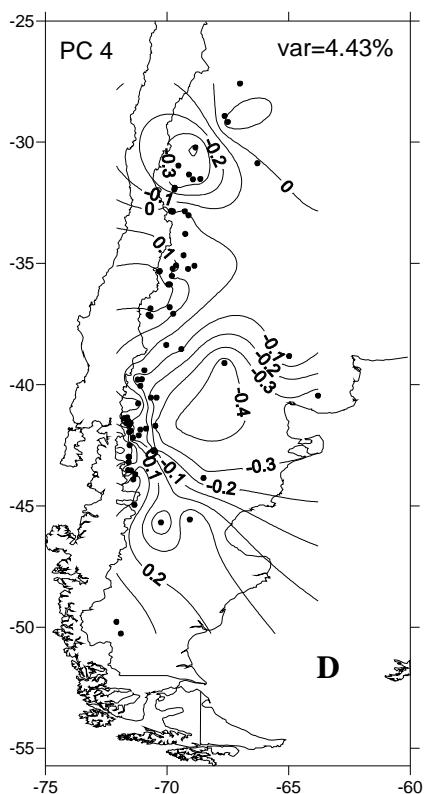
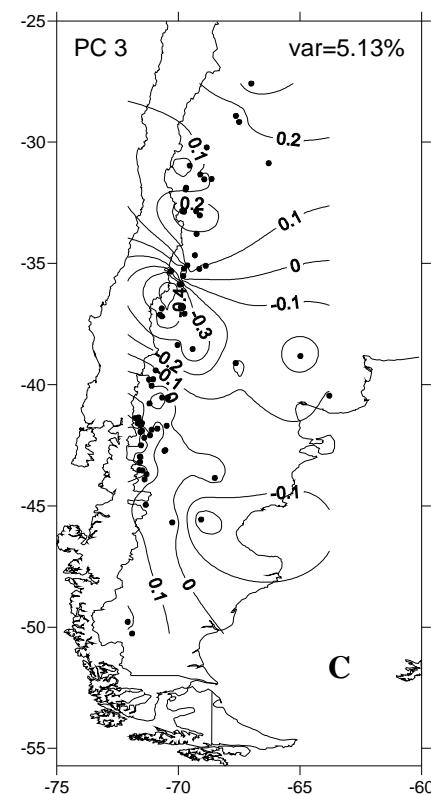
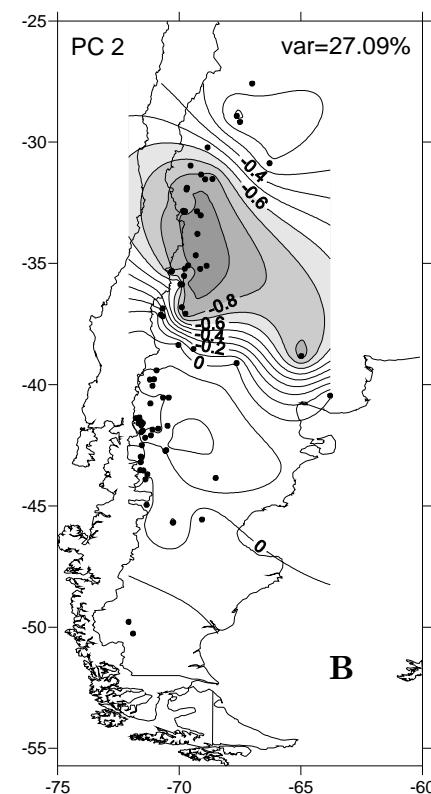
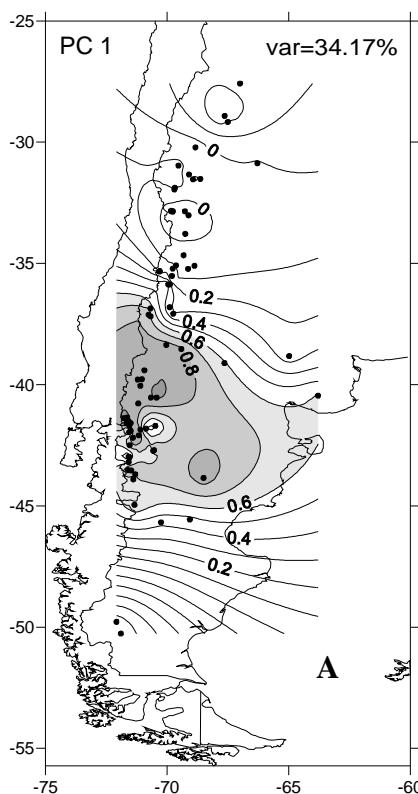
Localizar las distintas regiones a lo largo de la Cordillera, en las cuales el comportamiento de los caudales de los ríos cuyas nacientes se encuentran en las altas cumbres, muestran homogeneidades estadísticas, intentando definir zonas cuyas condiciones climáticas dan regímenes hídricos semejantes y caracterizando su variabilidad temporal.

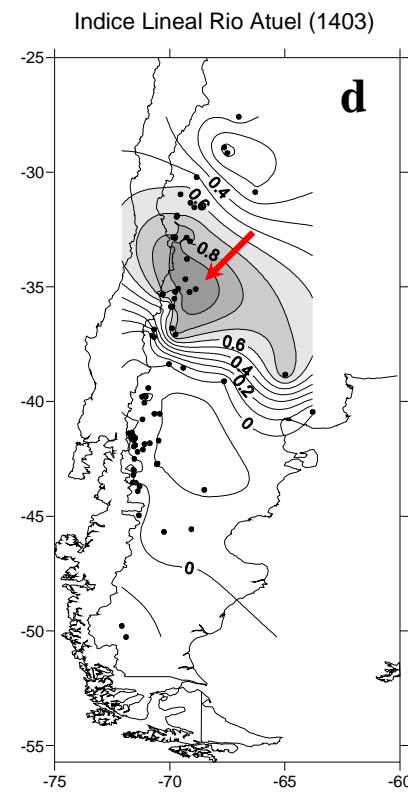
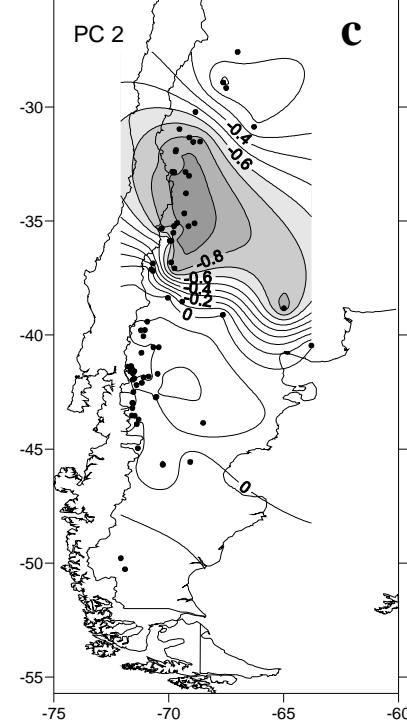
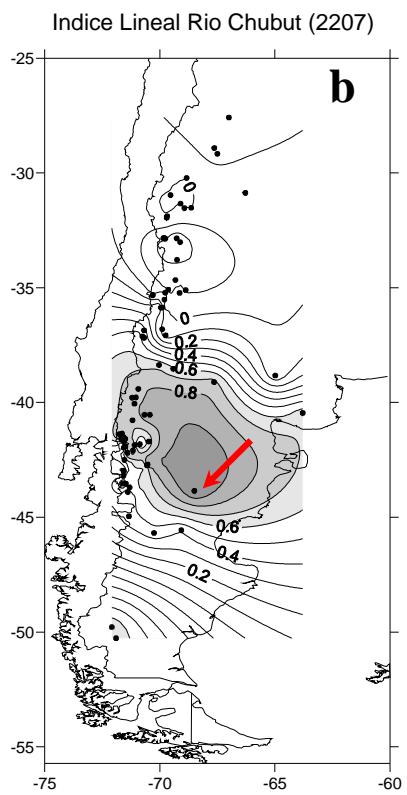
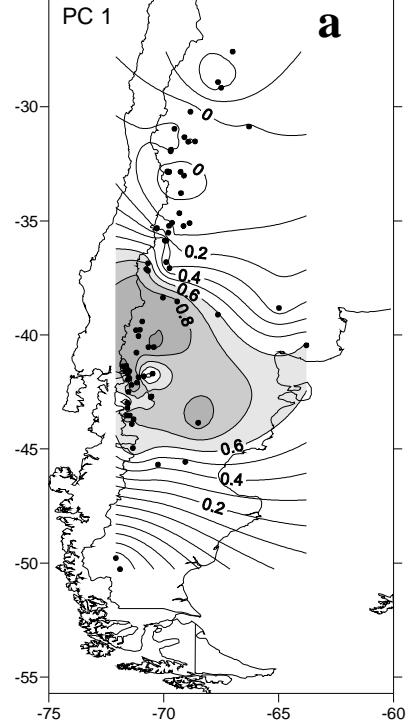
Análisis de Componentes Principales

Matriz de entrada
Correlaciones modificadas

$$\tilde{r} = \begin{cases} \delta(r) \frac{|r| - r_c}{1 - r_c} & \text{si } |r| > r_c \\ 0 & \text{si } |r| \leq r_c \end{cases}$$

r_c al 5% de significancia
 $\delta(r)$ función signo

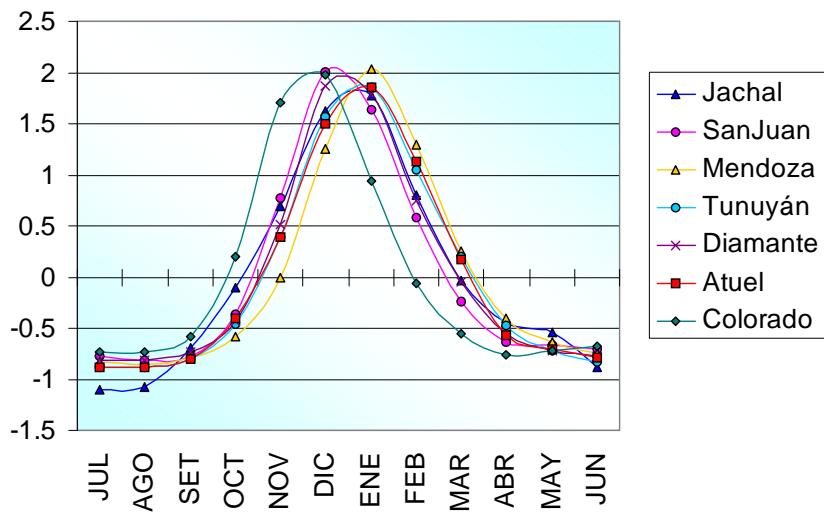




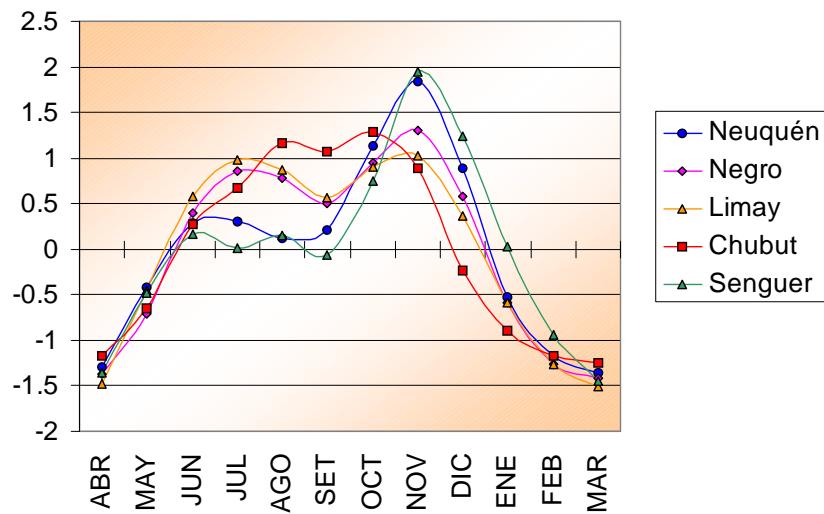
Estación Los Altares

Estación La Angostura

A Hidrogramas estandarizados



B Hidrogramas estandarizados



C Hidrogramas estandarizados

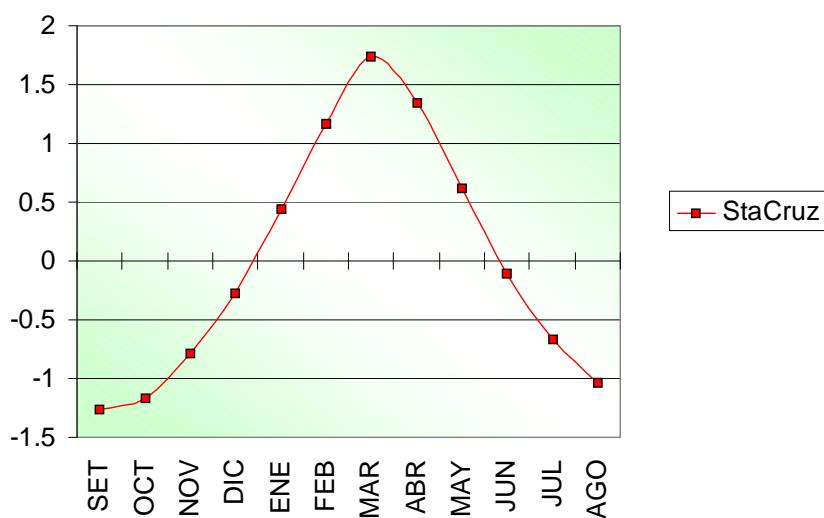
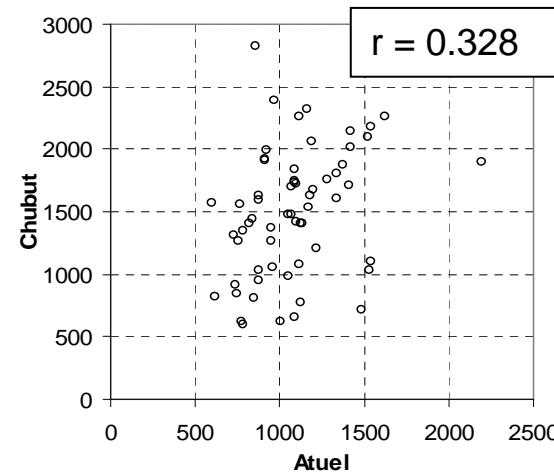
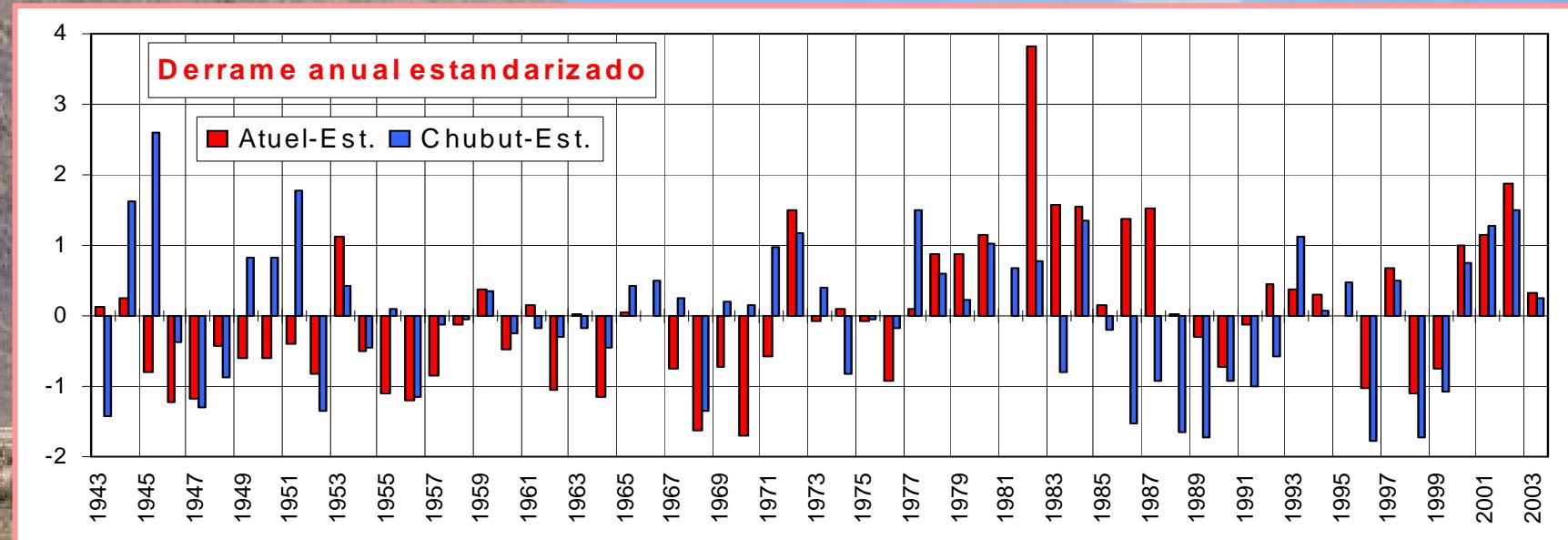


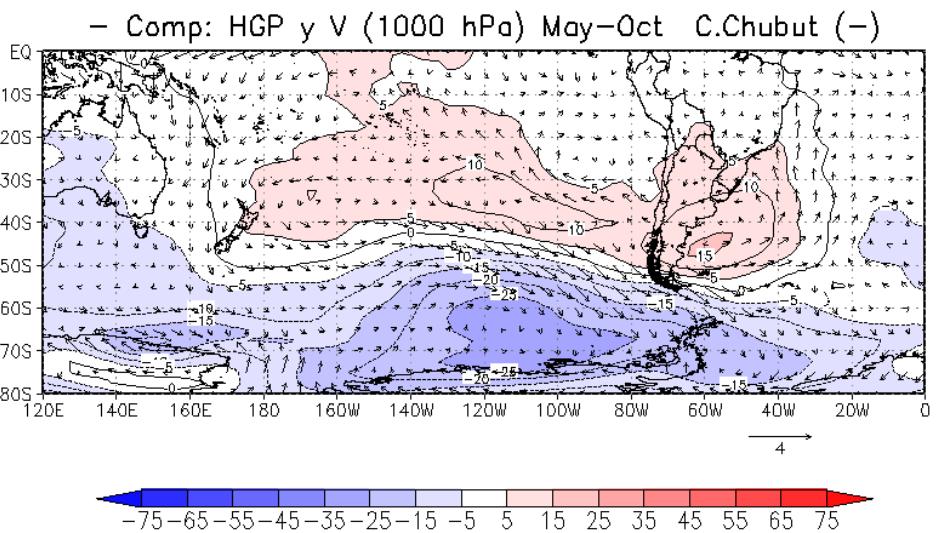
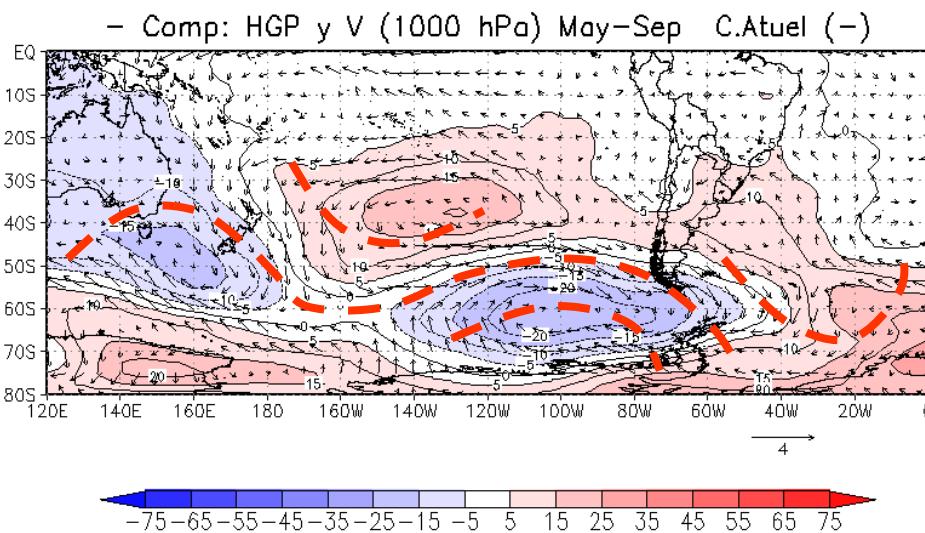
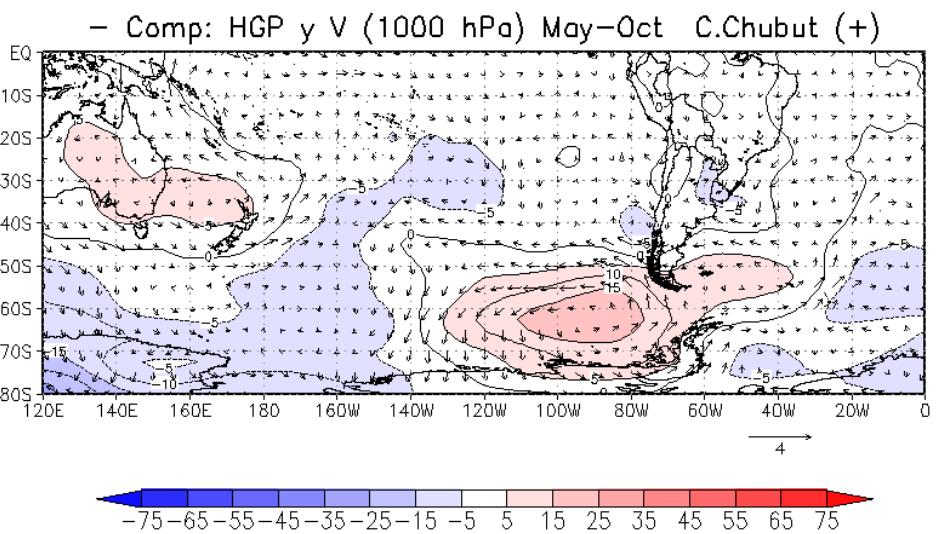
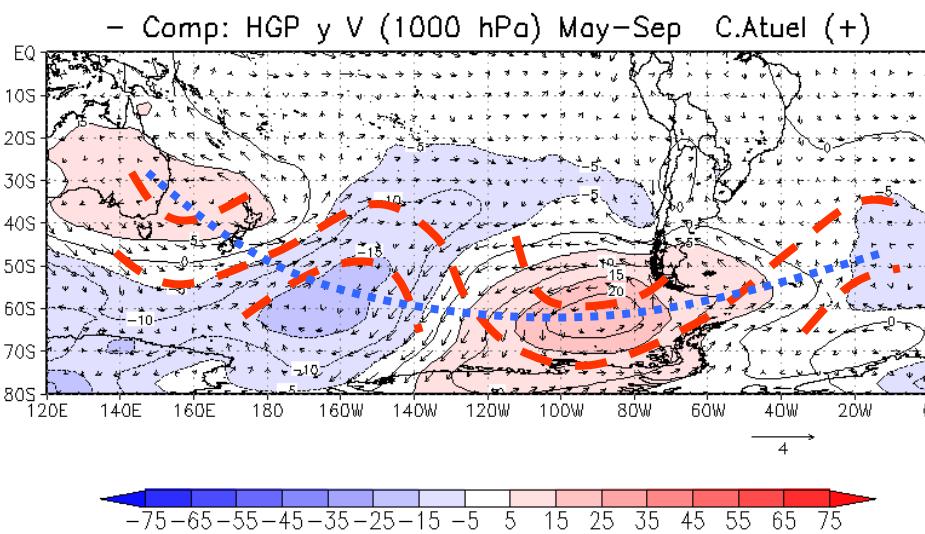
Diagrama de dispersión

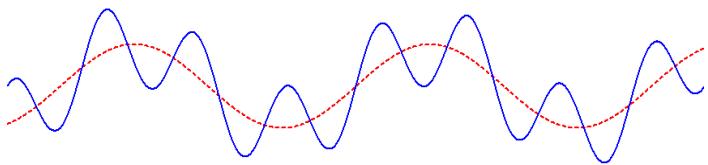


Análisis de Anomalías asociadas a extremos de Caudal en los ríos Atuel y Chubut

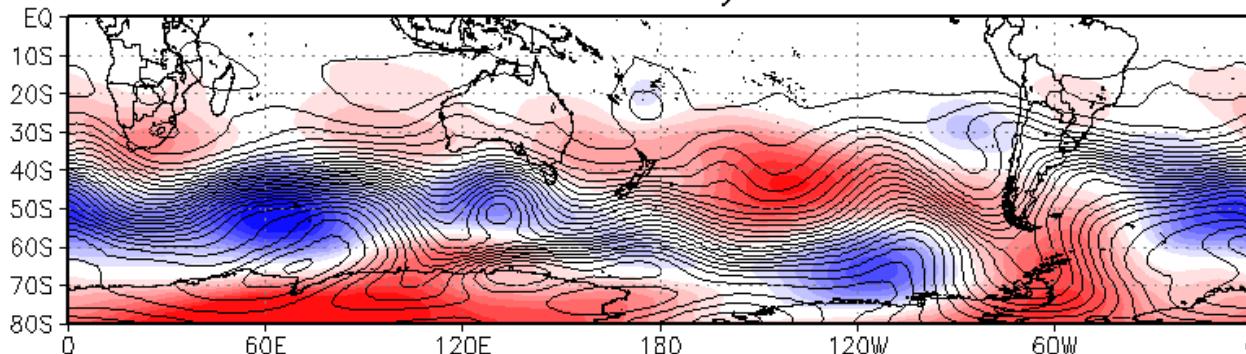


	Anomalía de caudal $> +\sigma$	Anomalía de caudal $< -\sigma$
Atuel	'72 '80 '82 '83 '84 '86 '87 '01 '02	'62 '64 '68 '70 '96 '98
Chubut	'72 '77 '80 '84 '93 '01 '02	'68 '86 '88 '89 '96 '98 '99

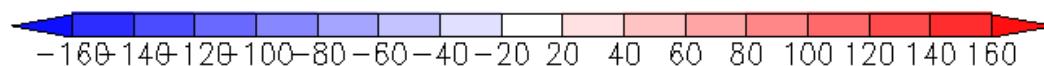




Comp: HGP (500 hPa) 13 days
Center: May 7



Ondas de Rossby cuasiestacionarias



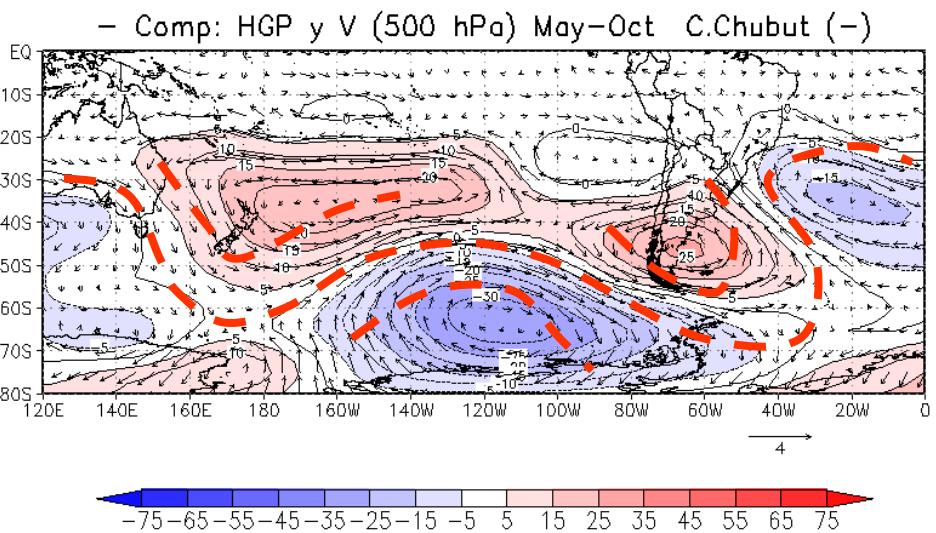
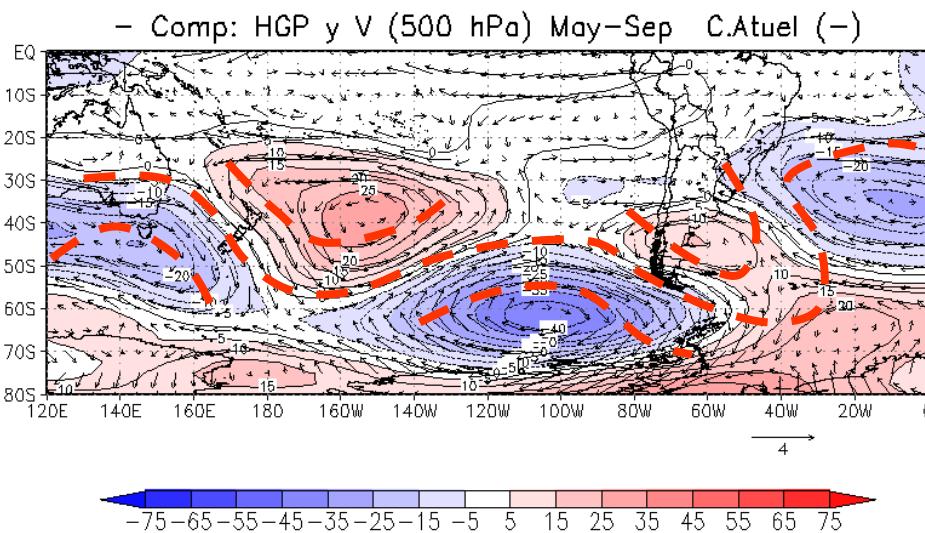
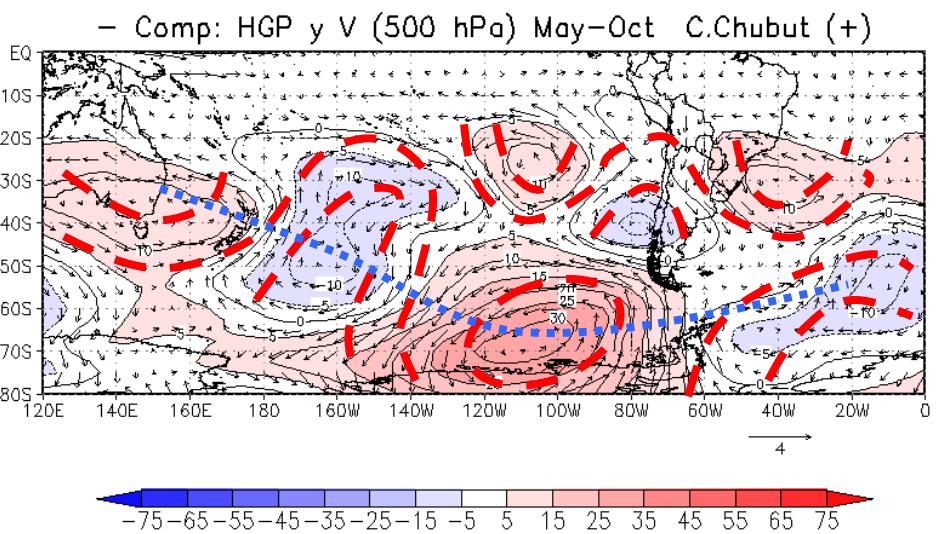
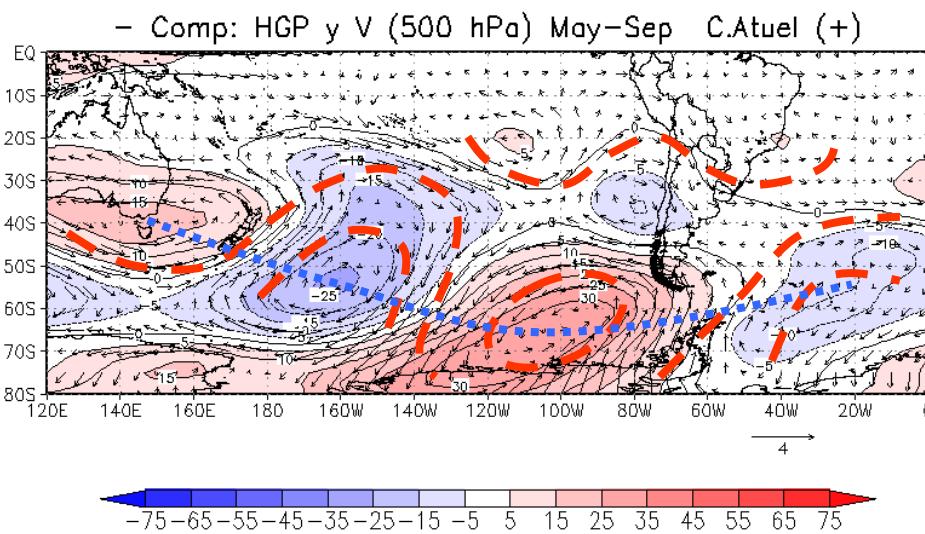
$$\frac{d(\xi + f)}{dt} = 0$$

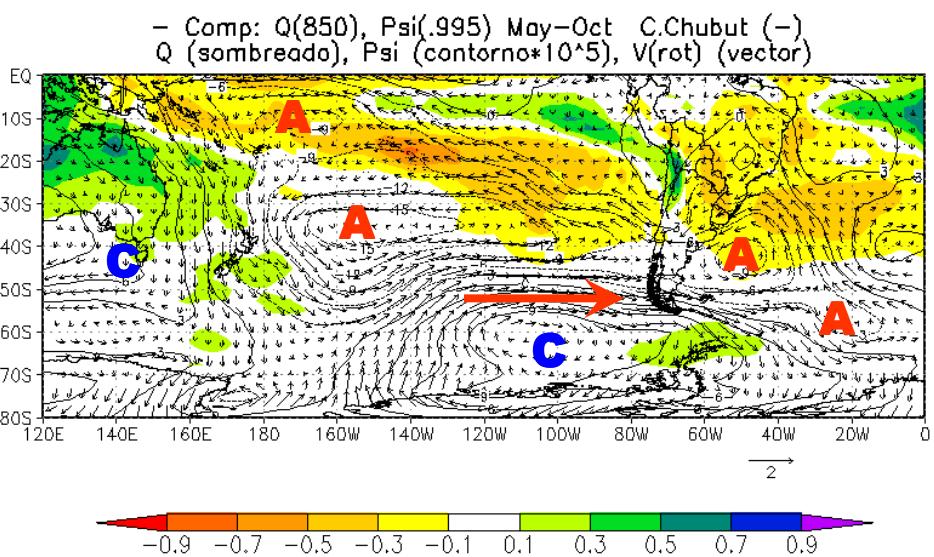
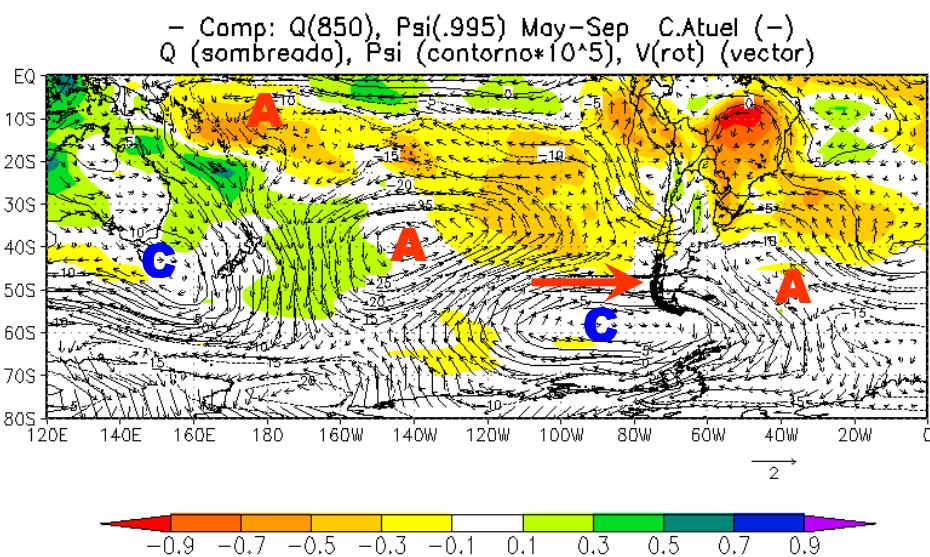
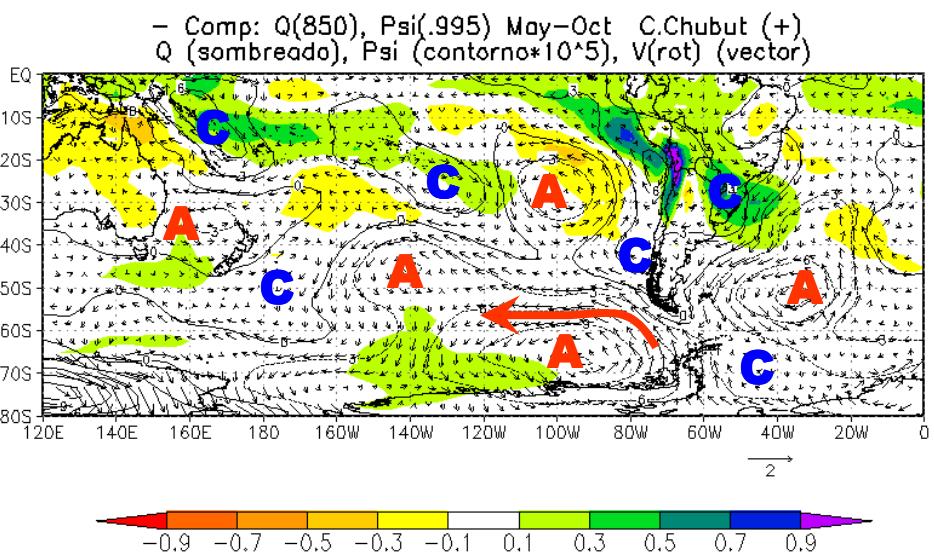
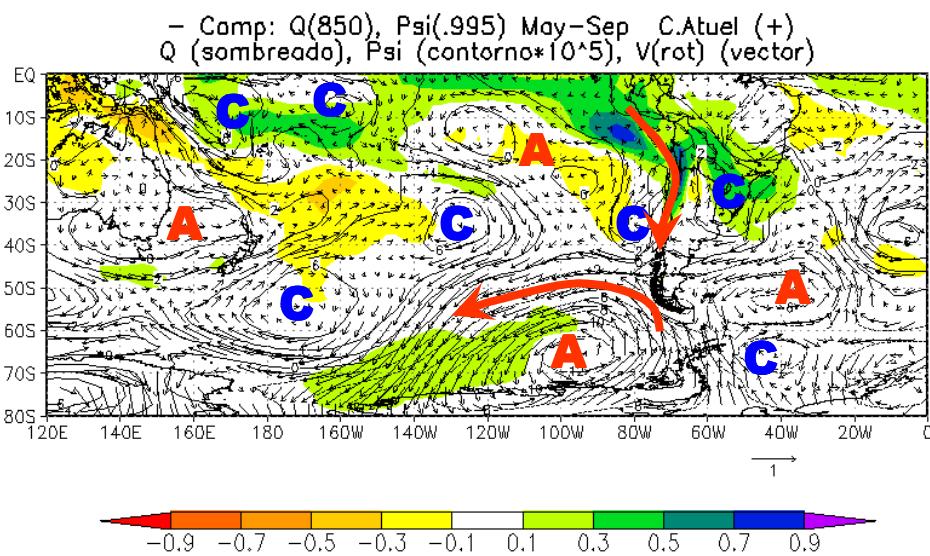
$$f = 2\Omega \sin(\phi)$$

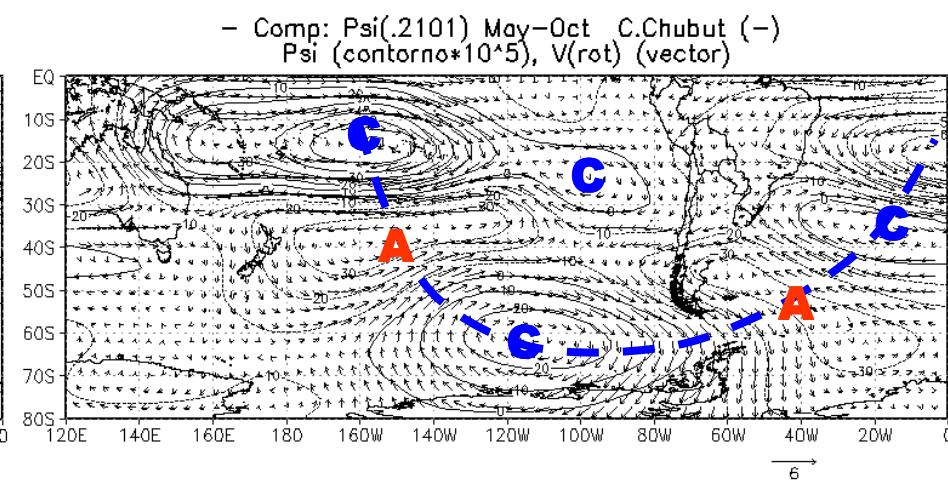
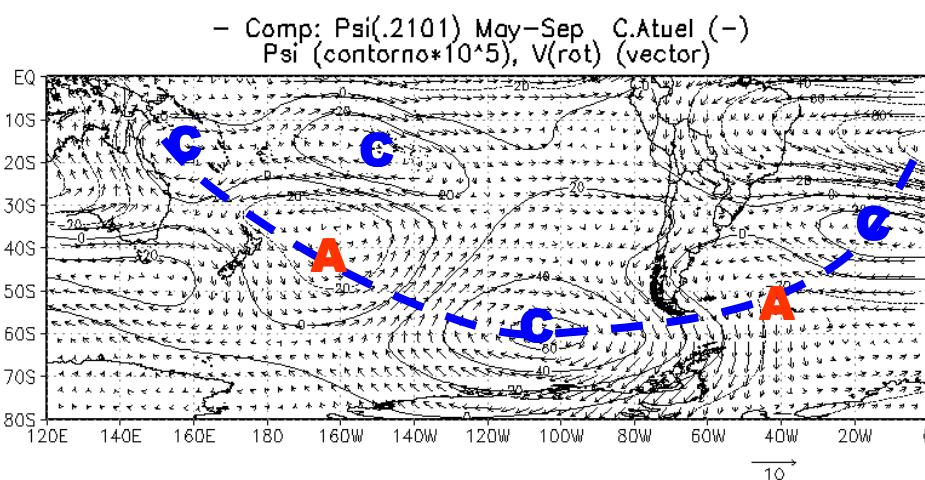
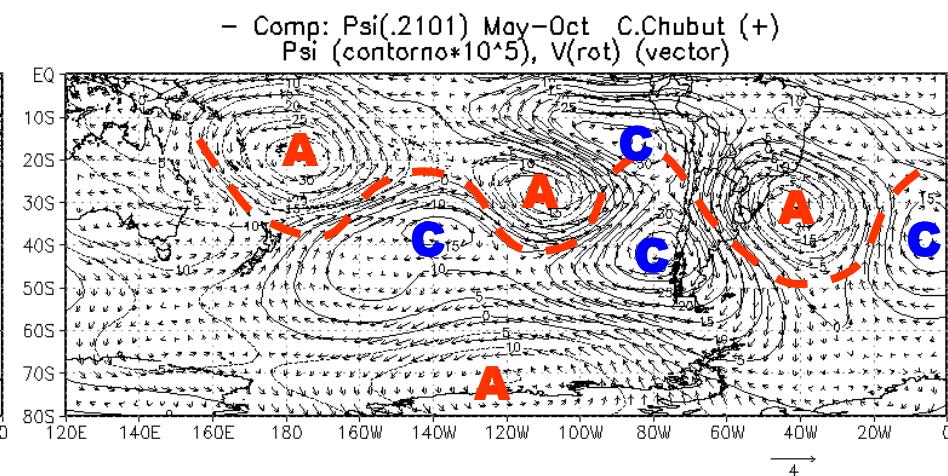
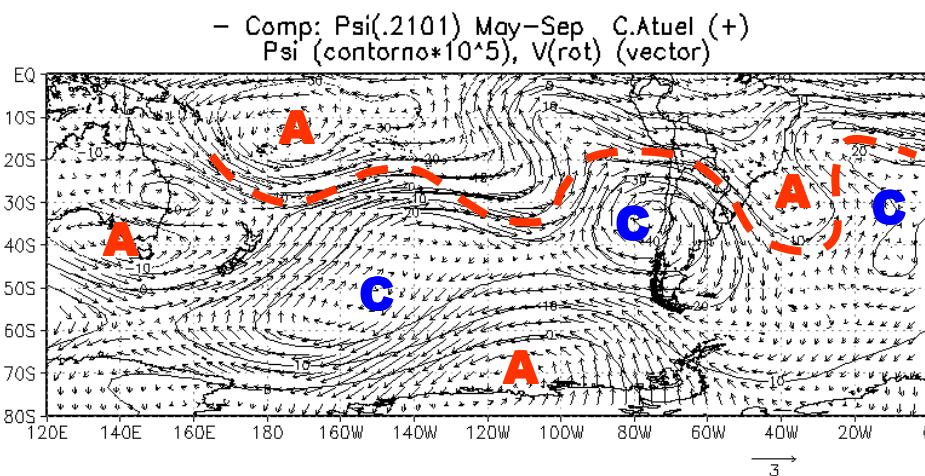
$$\xi = \frac{\partial u}{\partial y} - \frac{\partial v}{\partial x}$$

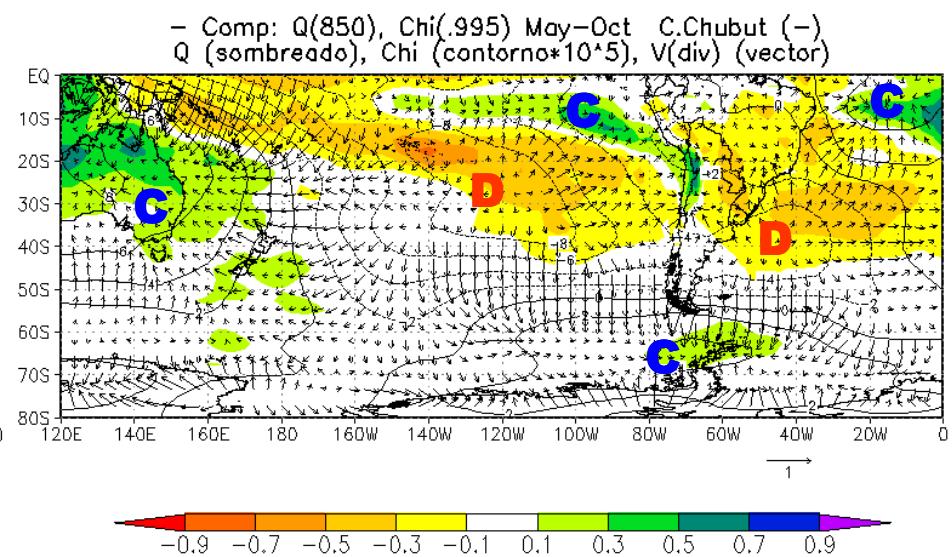
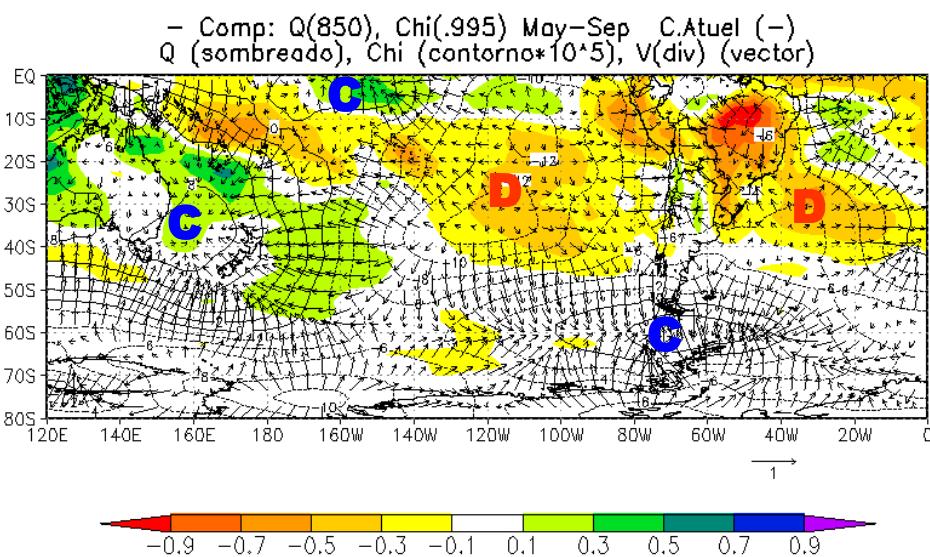
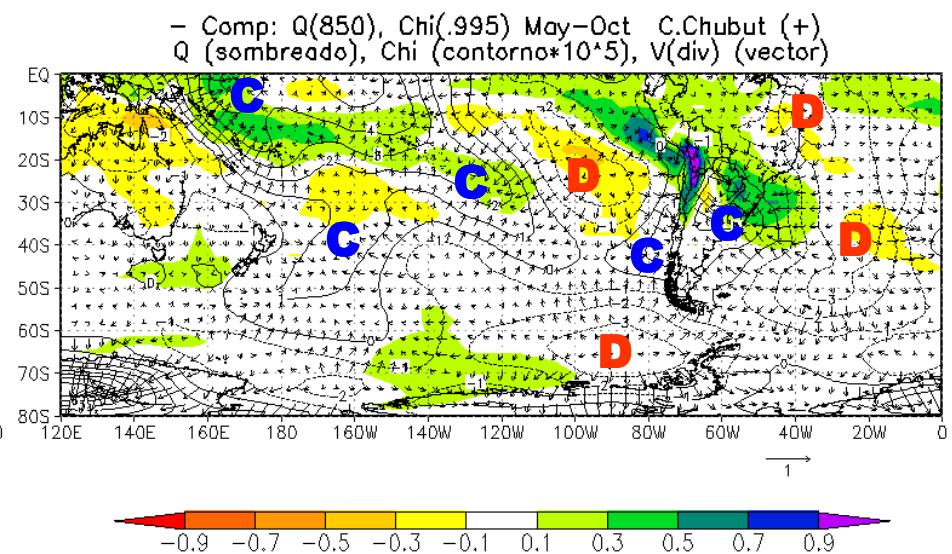
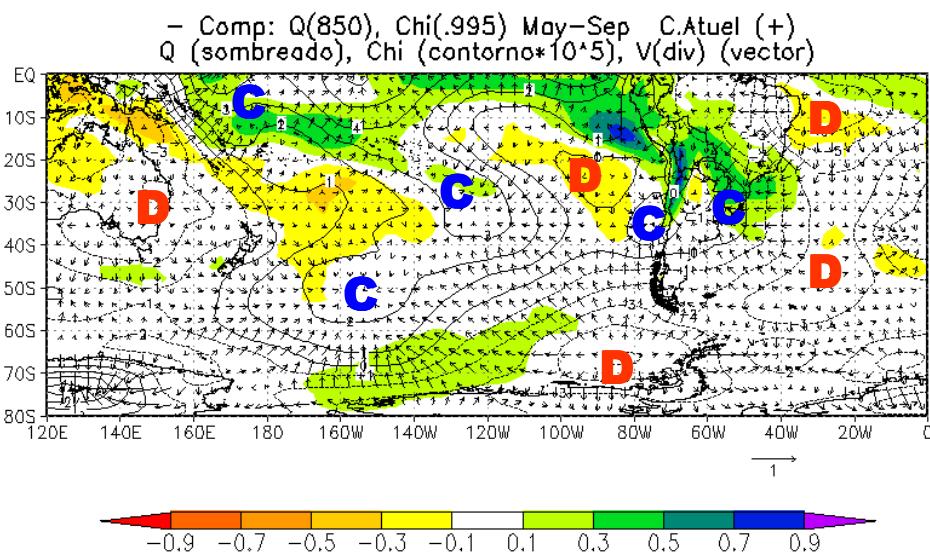
$$\frac{d\xi}{dt} = -\frac{df}{dt} = -2\Omega \cos(\phi) \frac{d\phi}{dt}$$

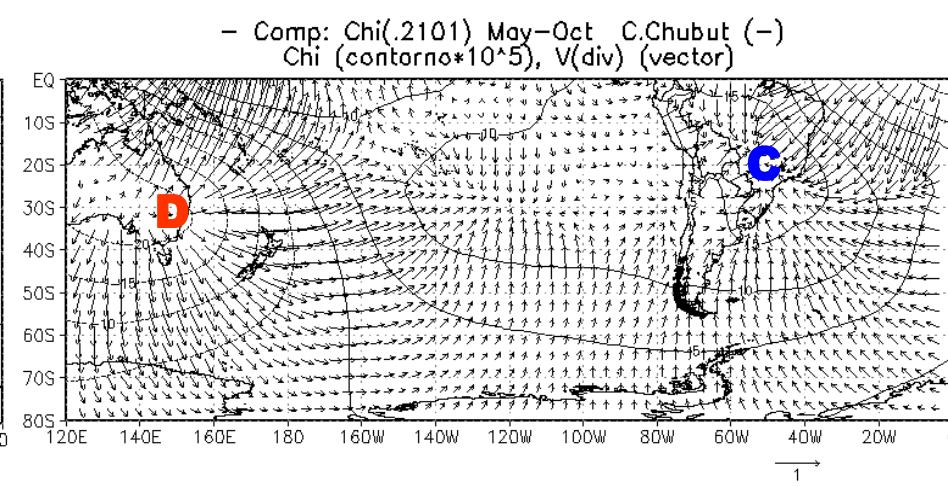
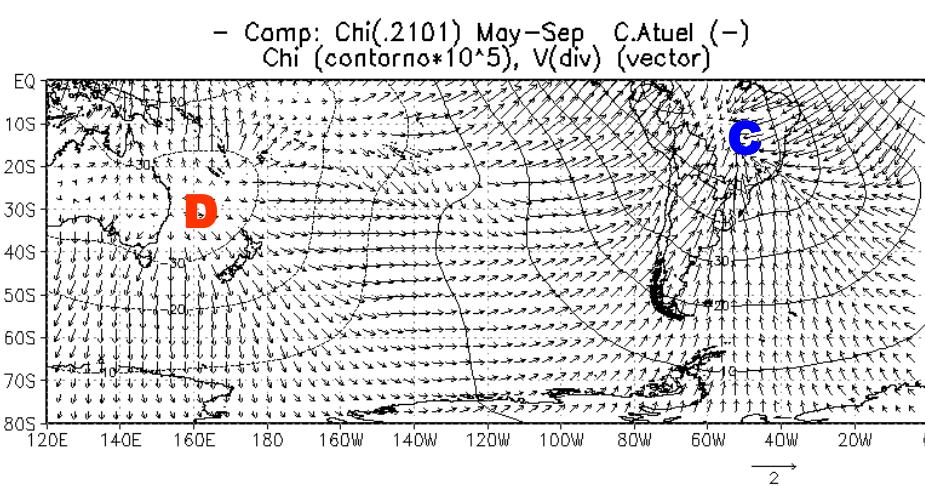
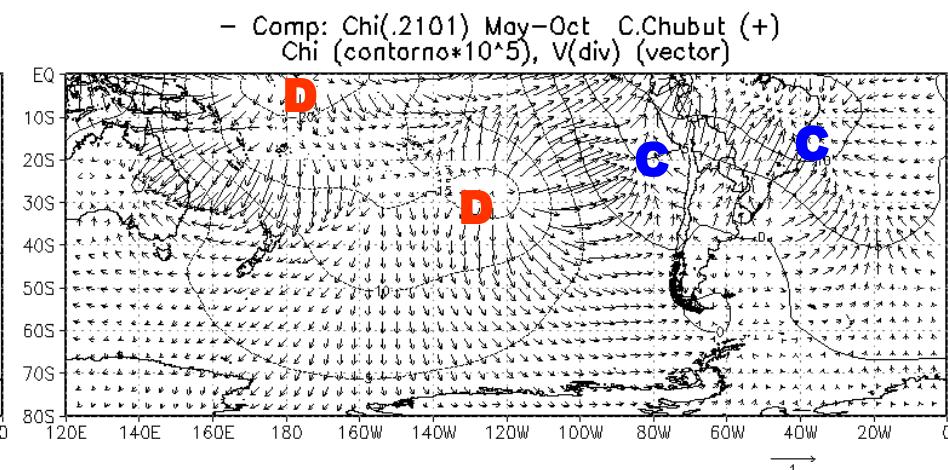
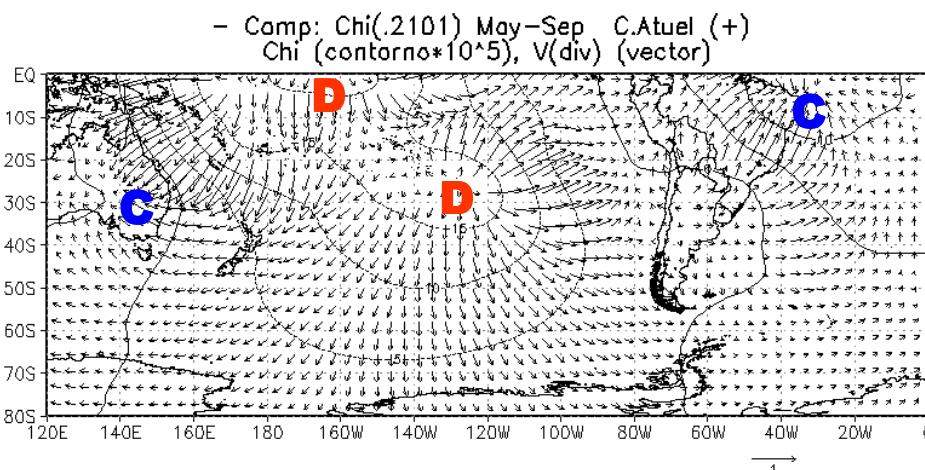
$$\Rightarrow \frac{d\xi}{dt} \propto -\frac{d\phi}{dt}$$











Propagación horizontal de ondas de Rossby estacionarias

Vectores de flujo de actividad de ondas
Plumb (1985)

$$\vec{F}_S = F_\lambda \cdot \hat{i} + F_\phi \cdot \hat{j}$$

$$F_\lambda = \frac{p}{2a^2 \cos(\phi)} \left[\left(\frac{\partial \psi'}{\partial \lambda} \right)^2 - \psi' \frac{\partial^2 \psi'}{\partial \lambda^2} \right] \quad F_\phi = \frac{p}{2a^2} \left[\frac{\partial \psi'}{\partial \lambda} \frac{\partial \psi'}{\partial \phi} - \psi' \frac{\partial^2 \psi'}{\partial \lambda \partial \phi} \right]$$

p = presión/1000 hPa,

ϕ es la latitud,

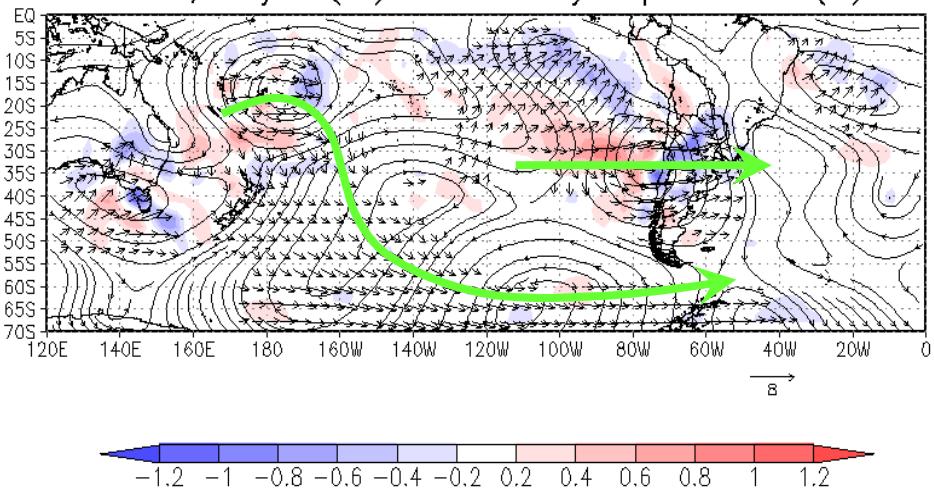
λ es la longitud,

a es el radio terrestre,

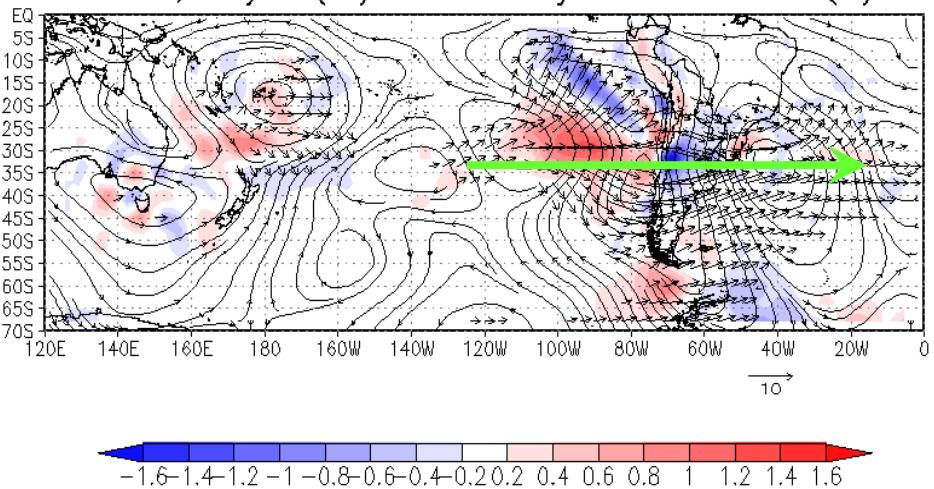
ψ es la función corriente y

las primas denotan la desviación respecto de la media zonal
(i.e. componentes transientes)

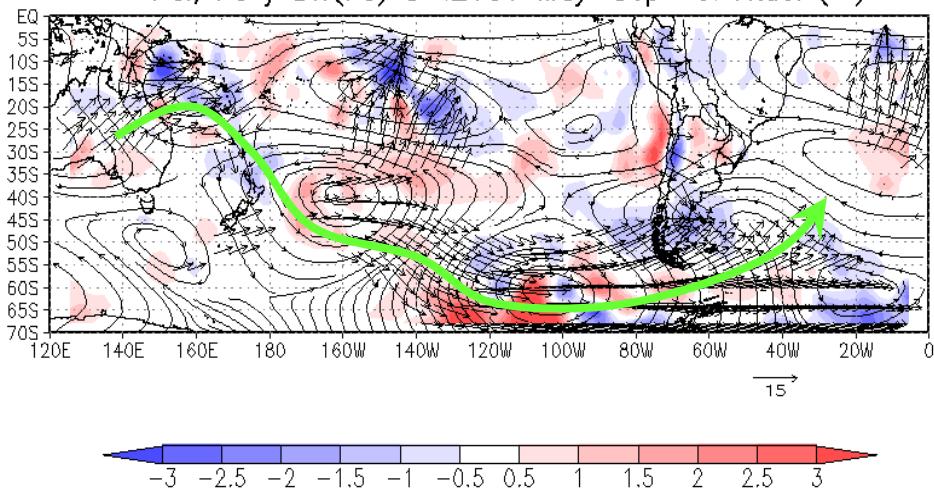
- Psi, Fs y Div(Fs) s=.2101 May-Sep C. Atuel (+)



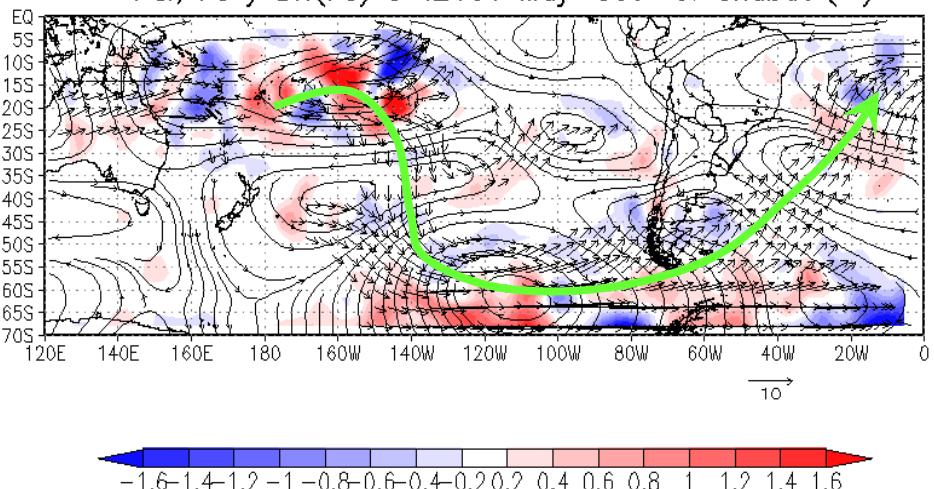
- Psi, Fs y Div(Fs) s=.2101 May-Oct C. Chubut (+)



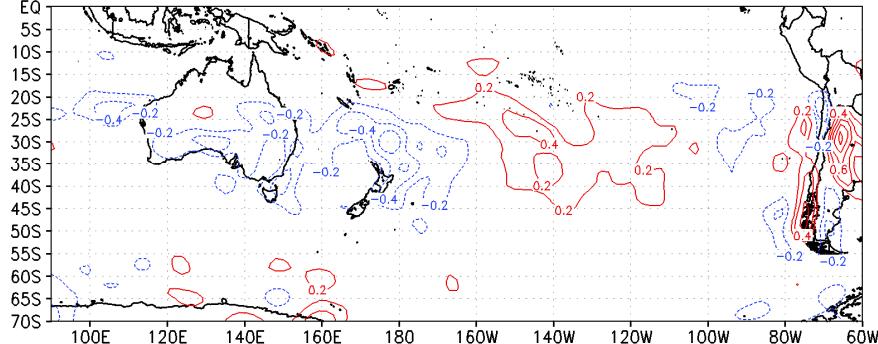
- Psi, Fs y Div(Fs) s=.2101 May-Sep C. Atuel (-)



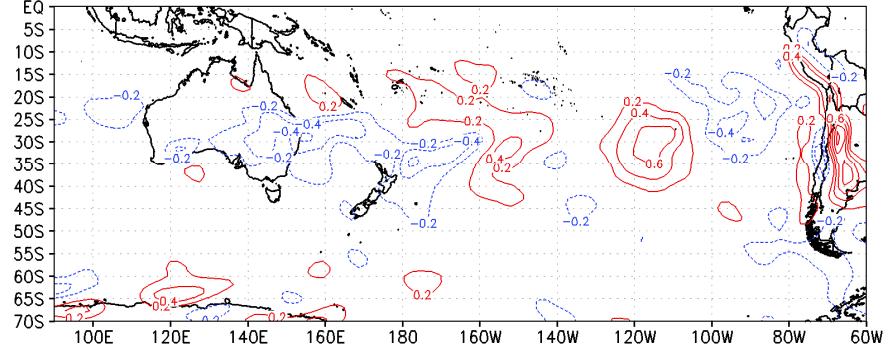
- Psi, Fs y Div(Fs) s=.2101 May-Oct C. Chubut (-)



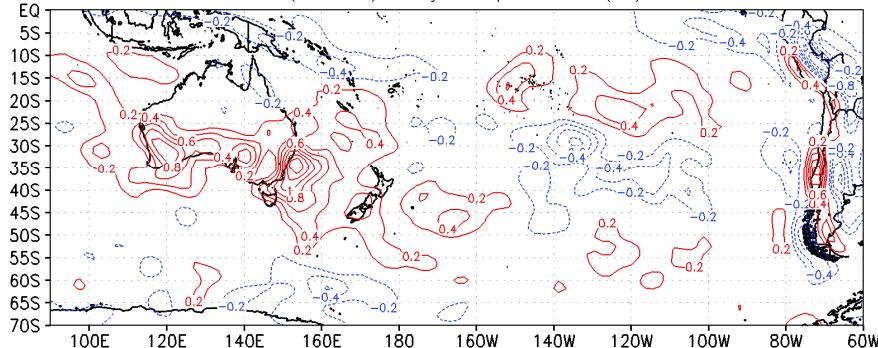
S(.2101) May–Sep Atuel (+)



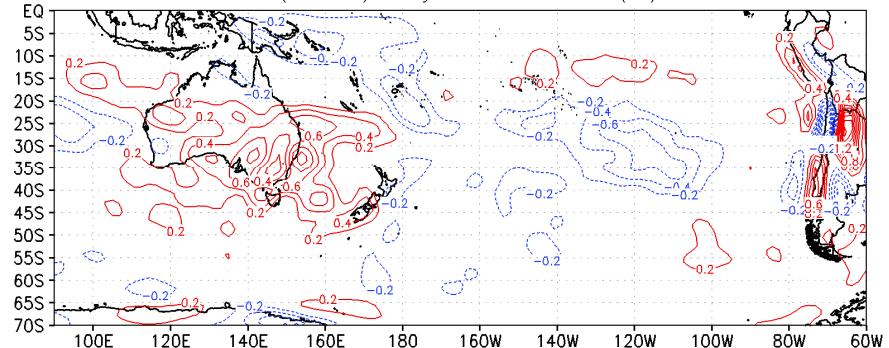
S(.2101) May–Oct Chubut (+)



S(.2101) May–Sep Atuel (-)



S(.2101) May–Oct Chubut (-)



$$\bar{S} = -\overline{f \nabla \cdot \mathbf{V}_\chi} - \overline{\xi \nabla \cdot \mathbf{V}_\chi} - \overline{\mathbf{V}_\chi \cdot \nabla (\xi + f)}$$

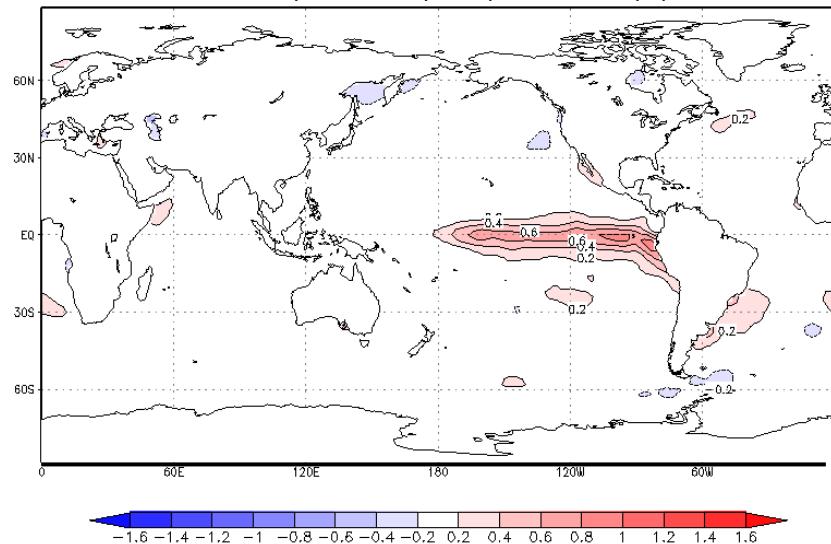


Temperatura Superficial del Mar y su relación con la variabilidad del Caudal de los ríos andinos

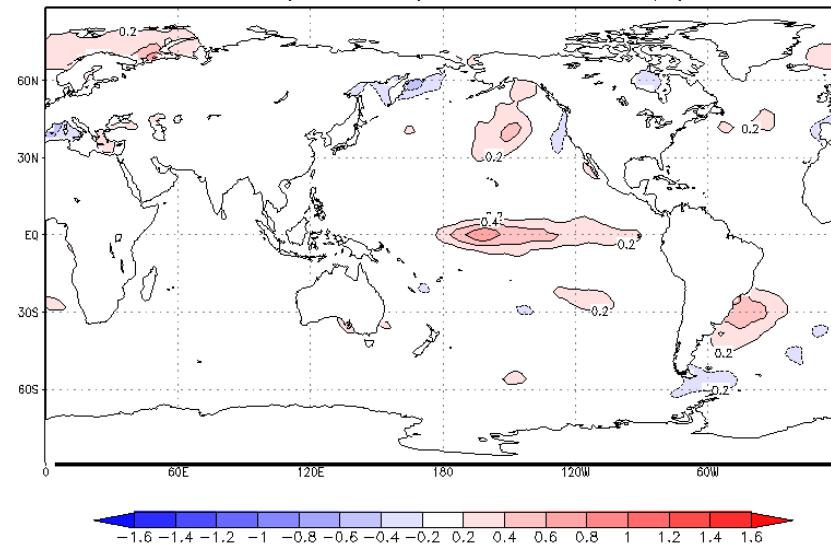
DATOS

Temperatura Superficial del Mar: Extended Reconstructed Sea Surface Temperature, ERSST, NOAA; construidos usando el más reciente conjunto de datos de 1SM del COADS (Comprehensive Ocean-Atmosphere Data Set) y mejorados mediante técnicas estadísticas que permiten una reconstrucción estable utilizando valores escasos. Período 1958-2004

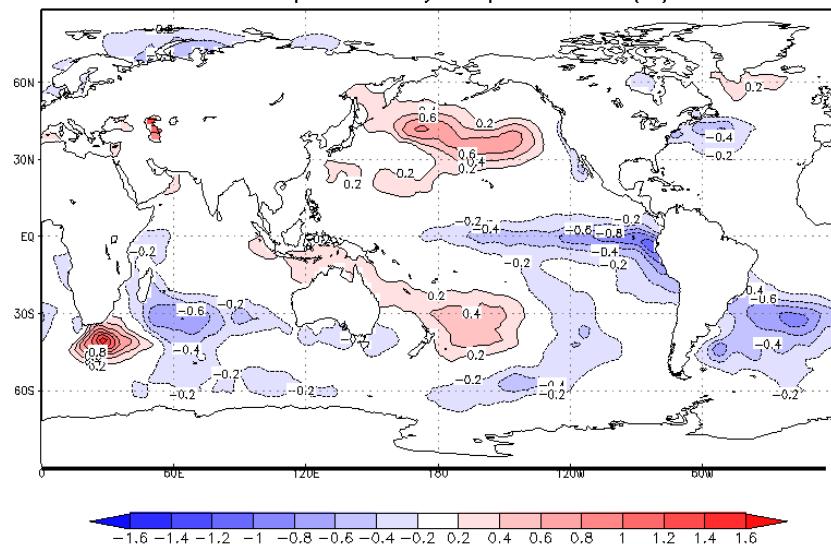
- Comp: SST May-Sep C.Atuel (+)



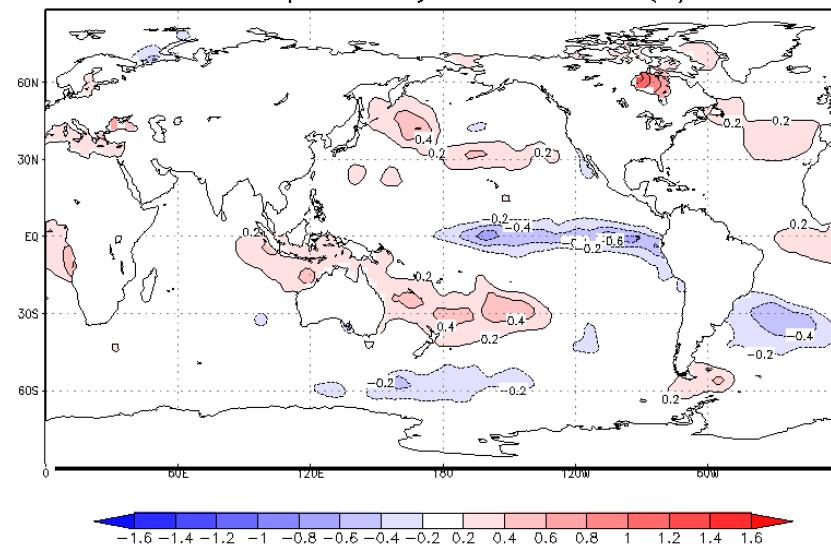
- Comp: SST May-Oct C.Chubut (+)



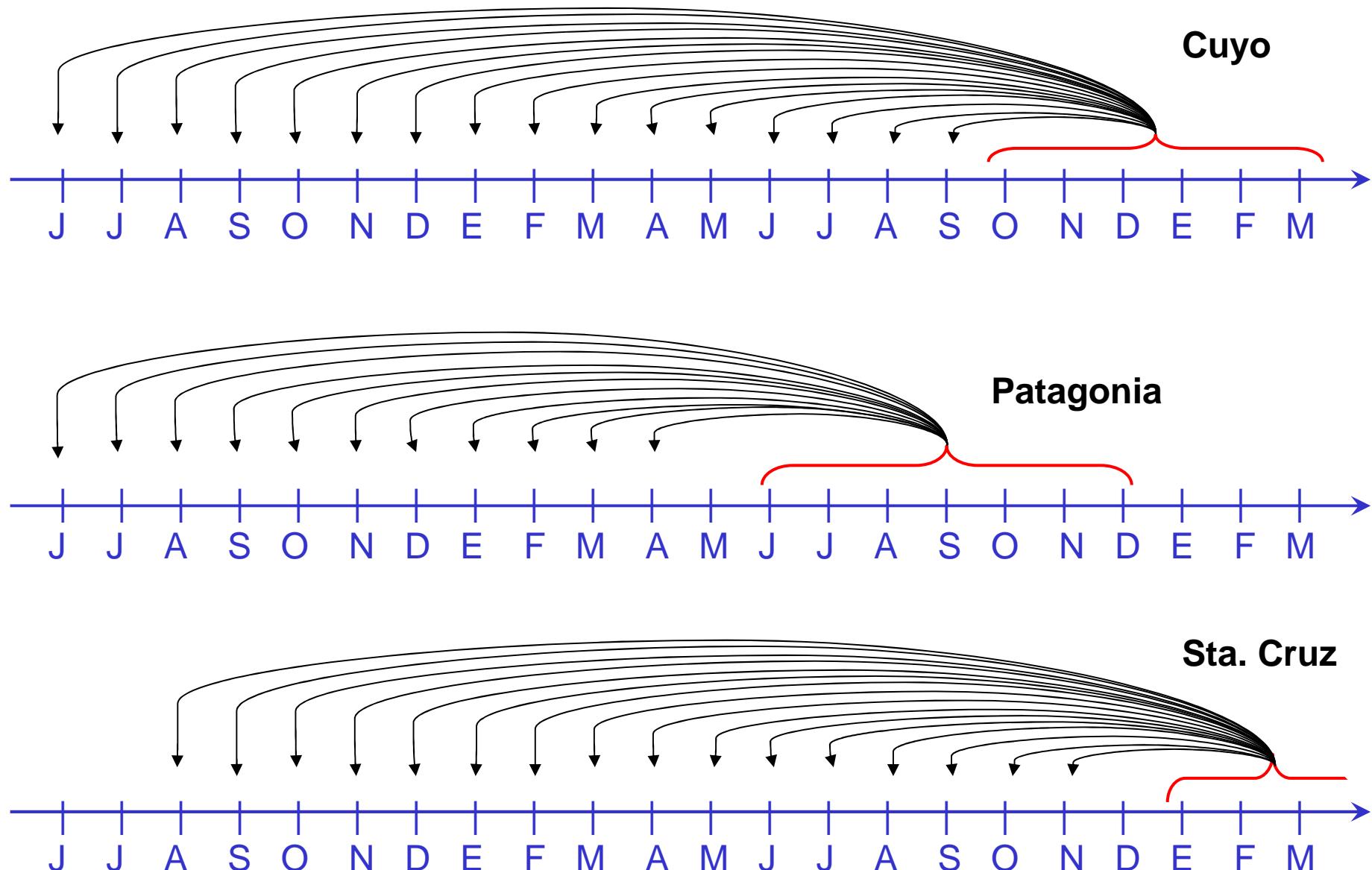
- Comp: SST May-Sep C.Atuel (-)



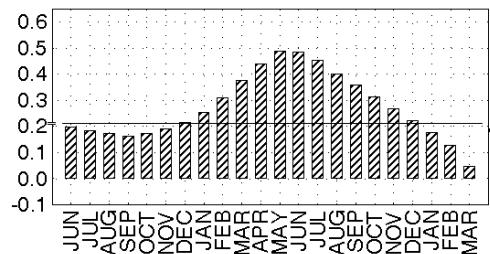
- Comp: SST May-Oct C.Chubut (-)



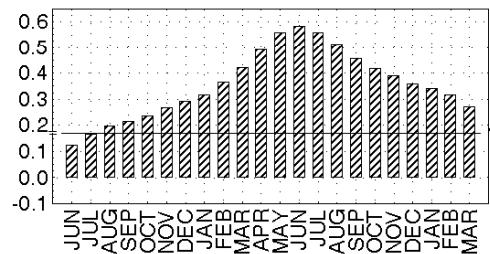
Funciones de correlación: Caudal estacional vs. Niño 3+4



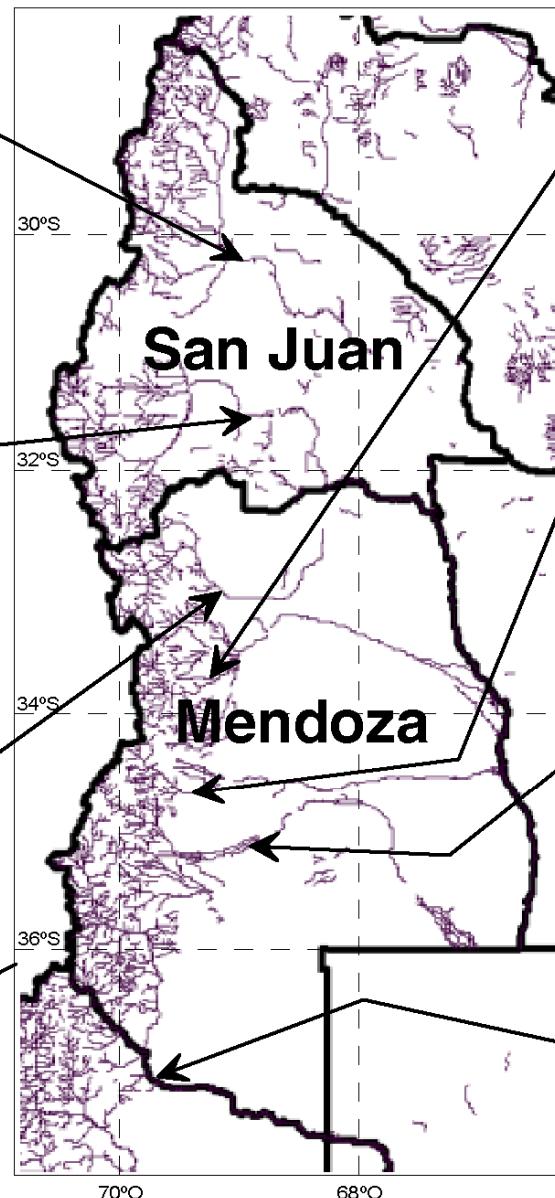
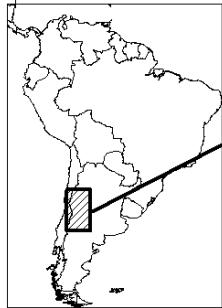
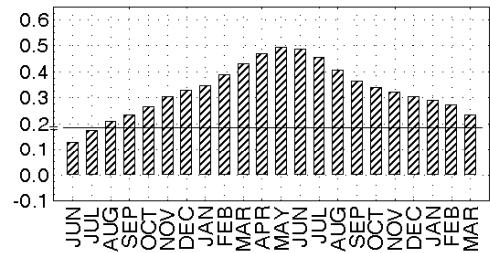
Jachal



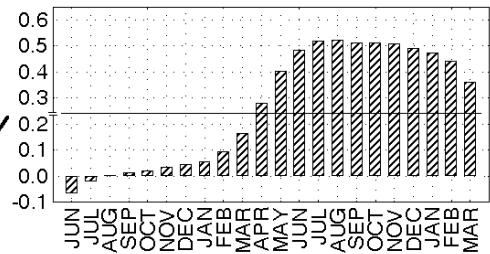
San Juan



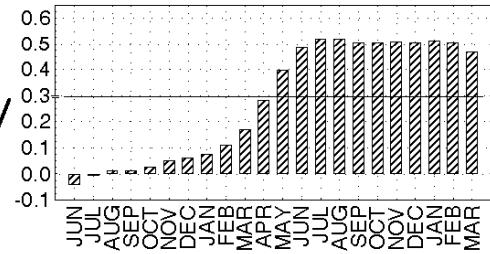
Mendoza



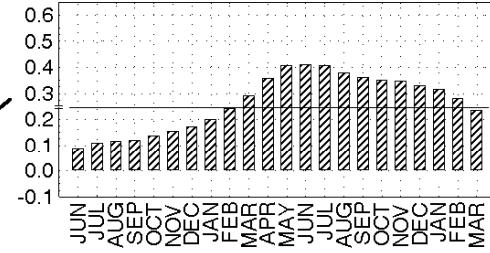
Tunuyán



Diamante



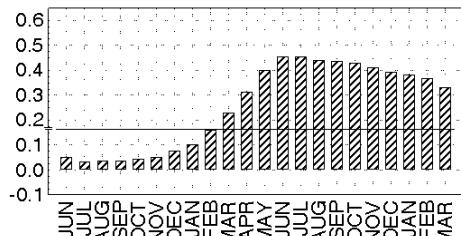
Atuel



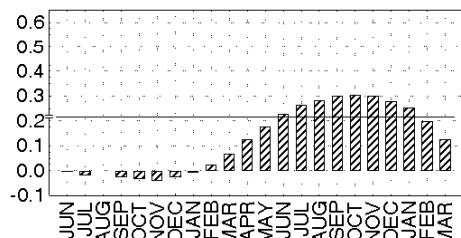
Colorado



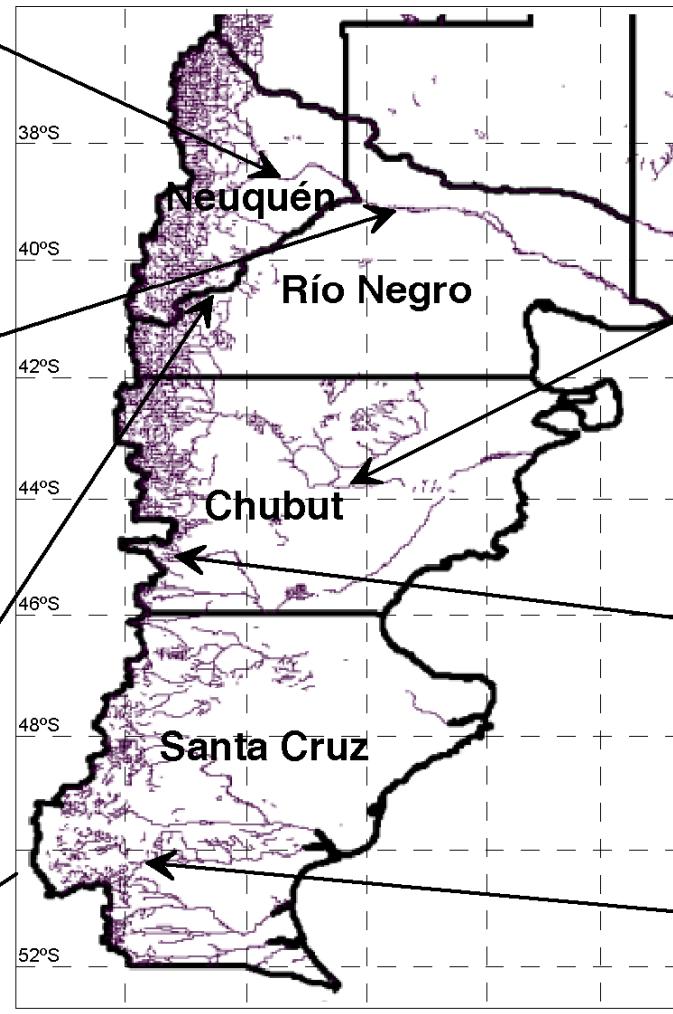
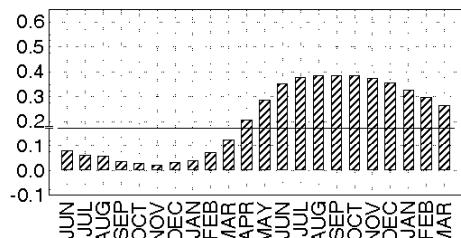
Neuquén



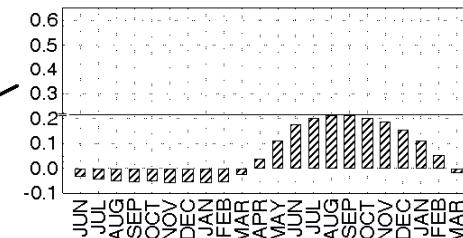
Negro



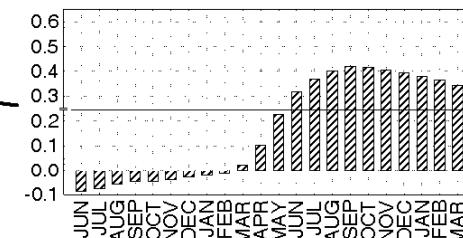
Limay



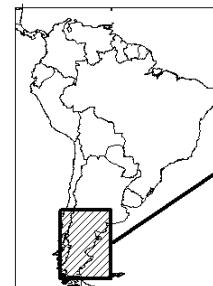
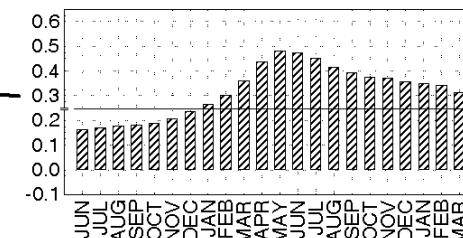
Chubut



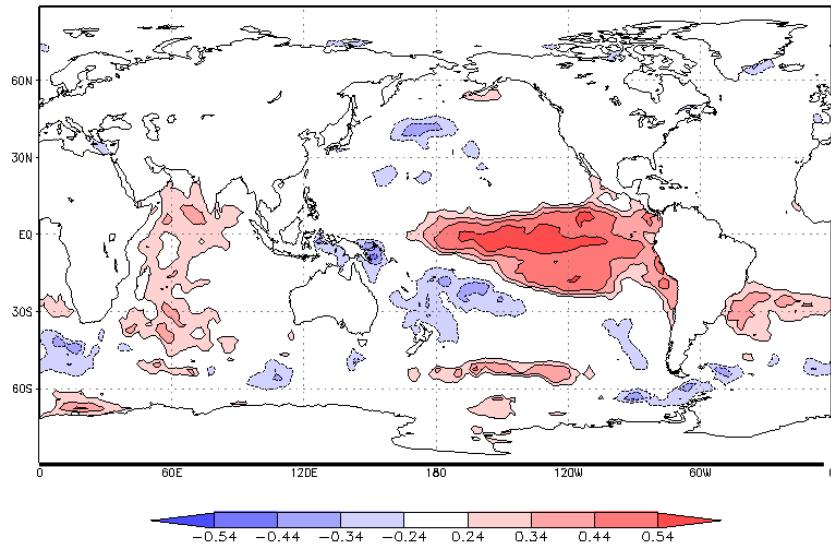
Senguer



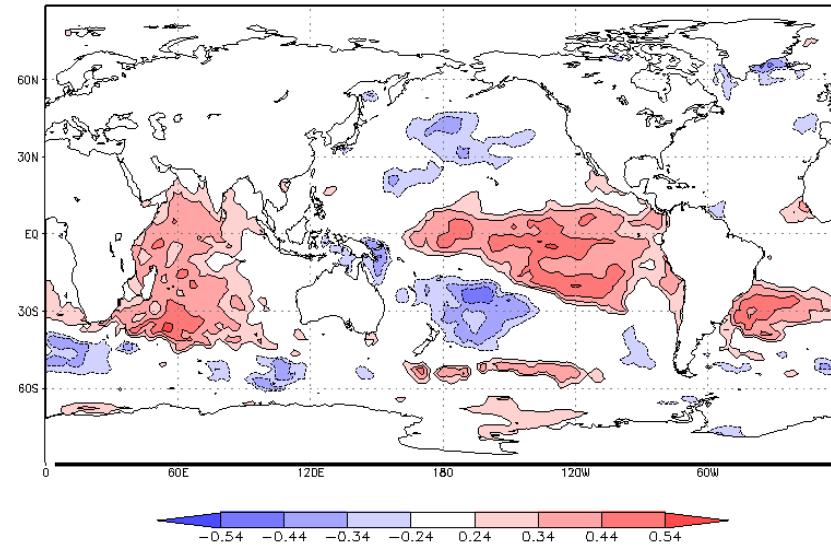
Santa Cruz



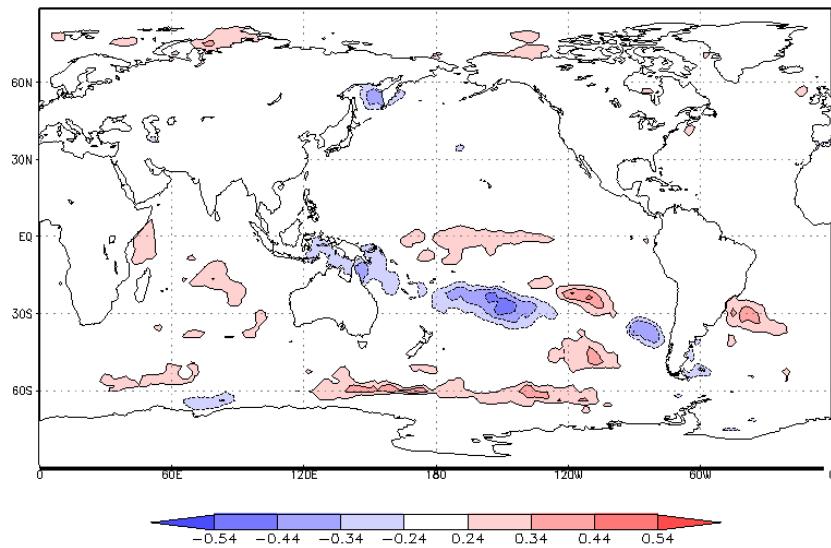
Corr: SST May–Ago vs. C San Juan Nov–Feb



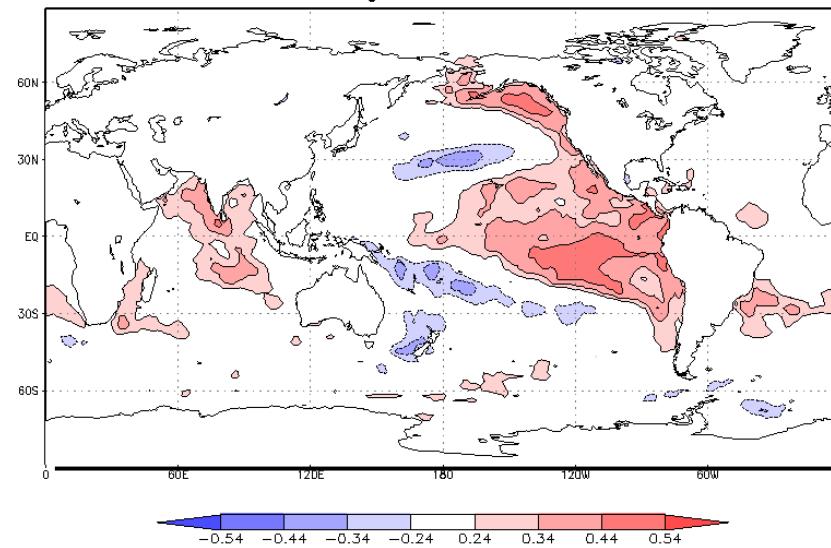
Corr: SST May–Sep vs. C Atuel Nov–Mar

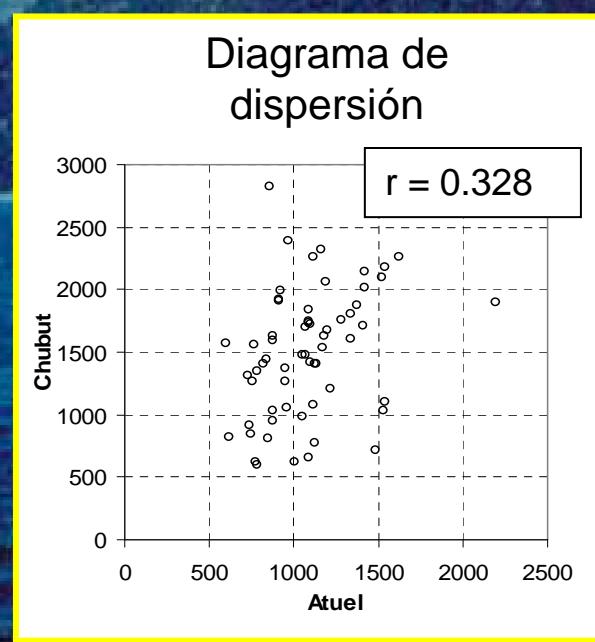
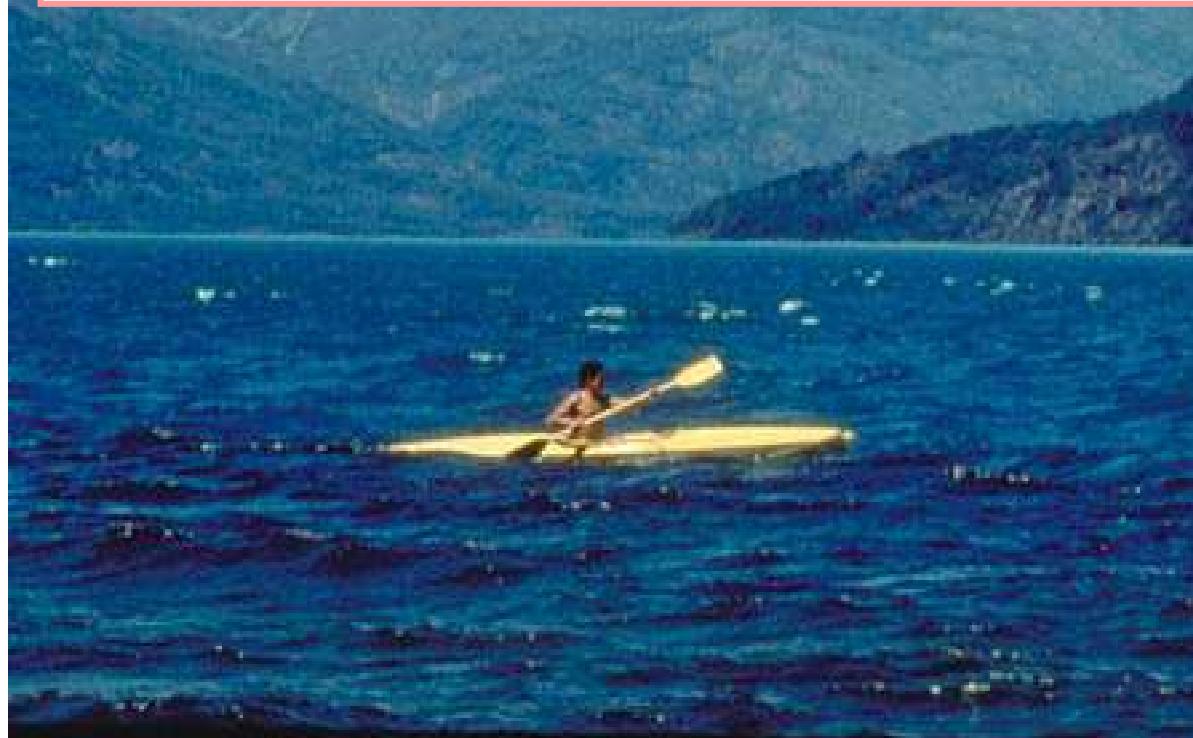
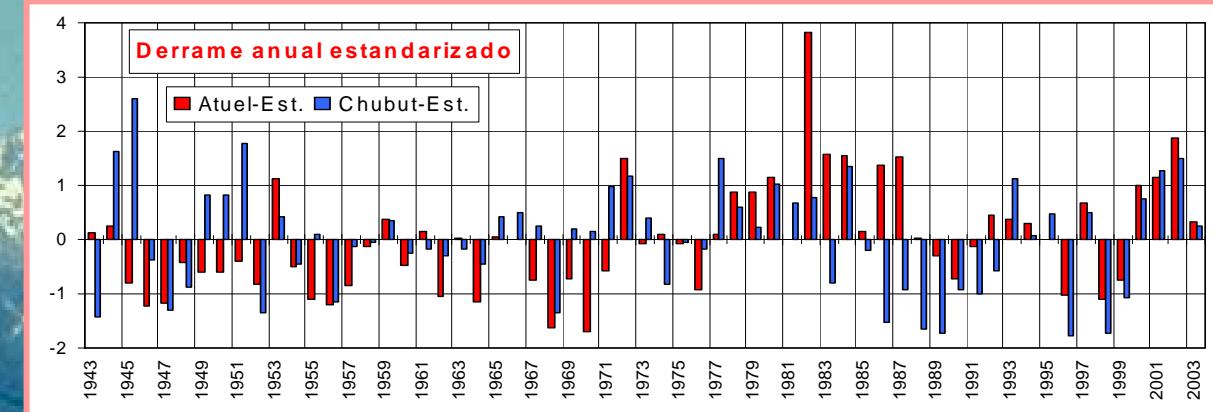


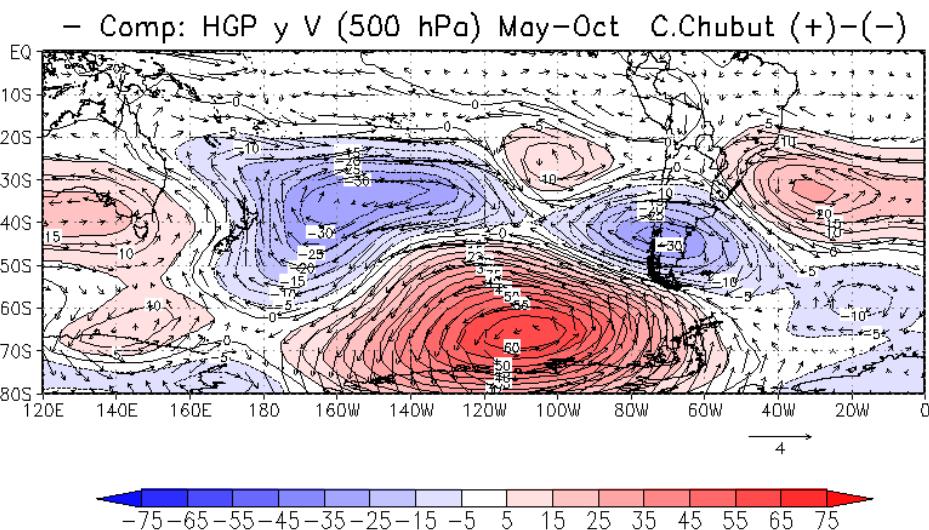
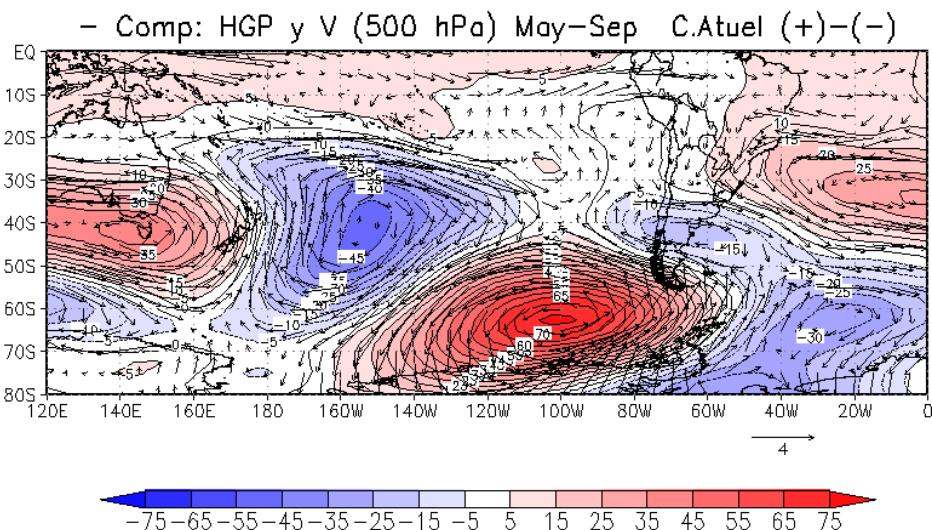
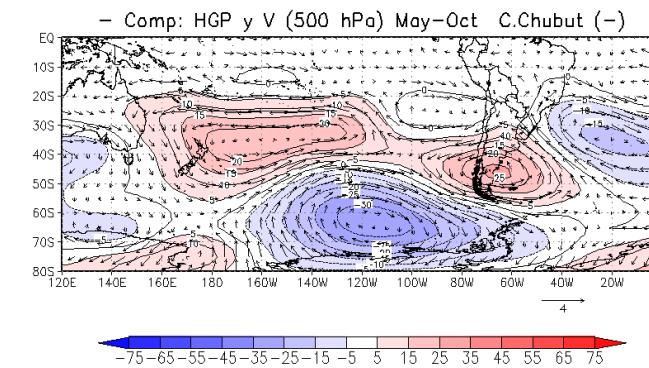
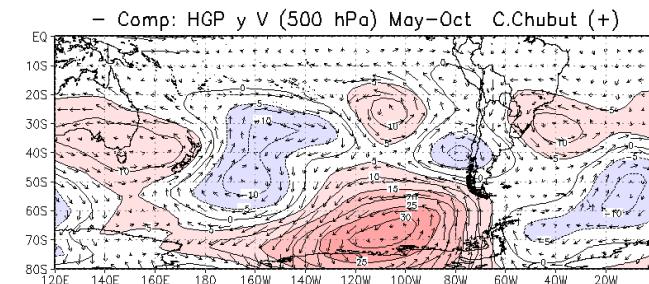
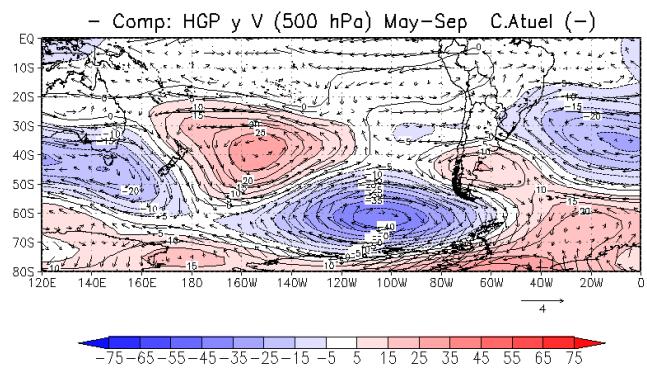
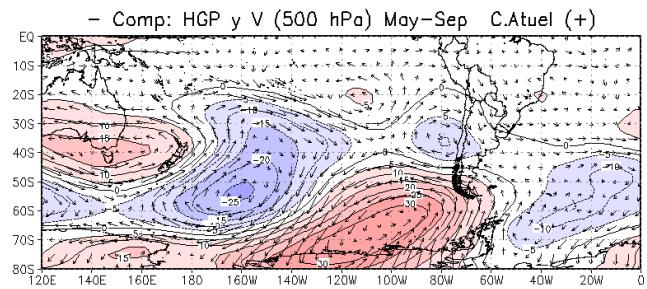
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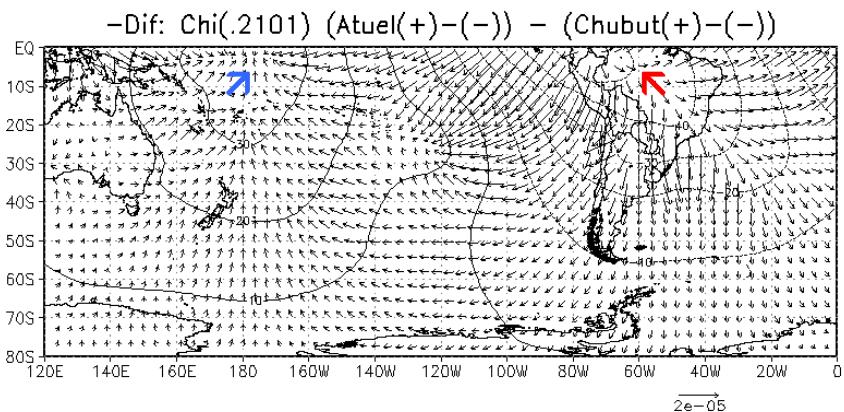
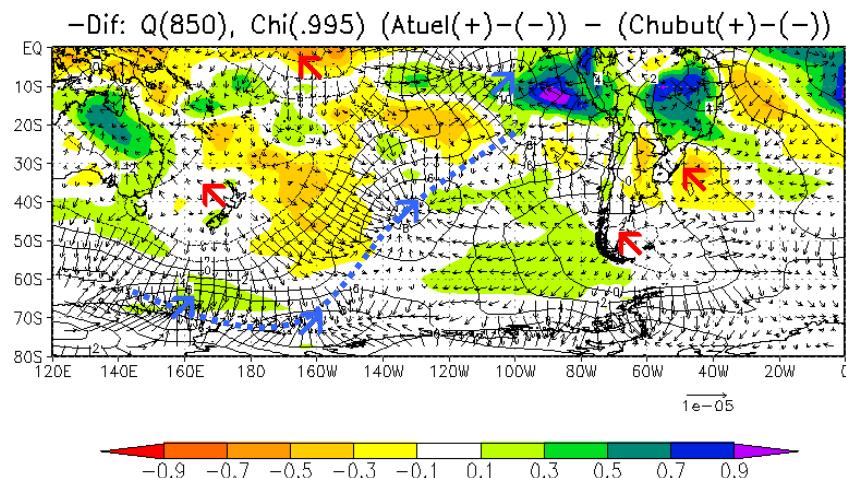
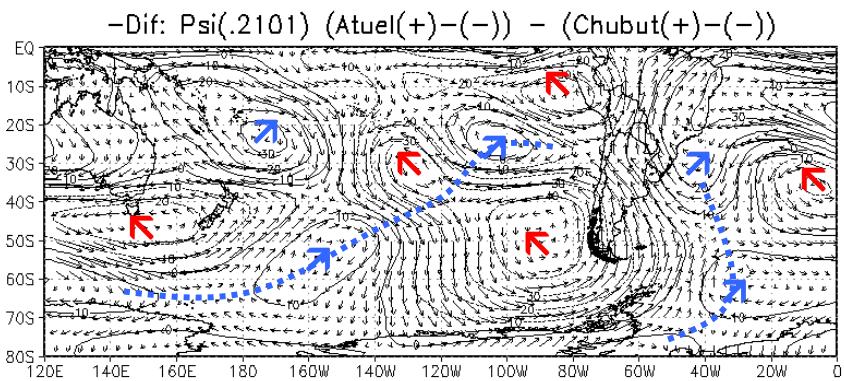
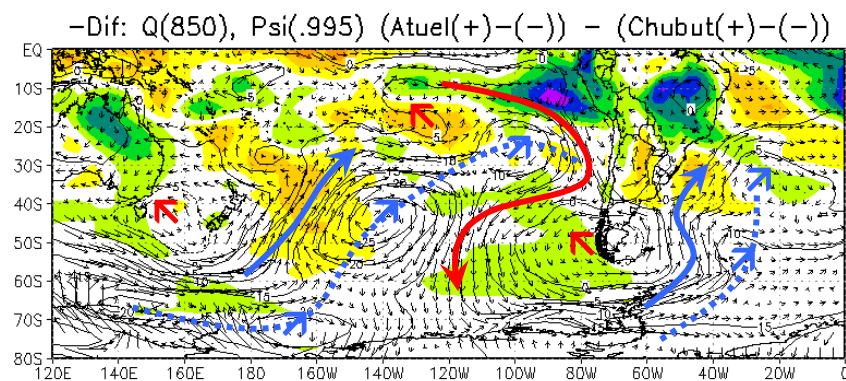
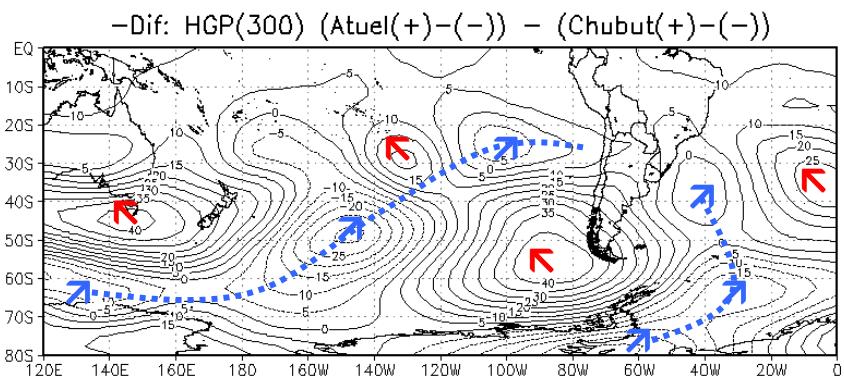
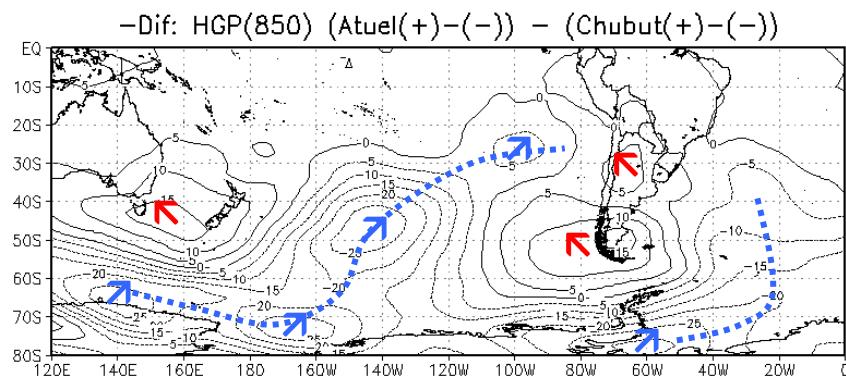


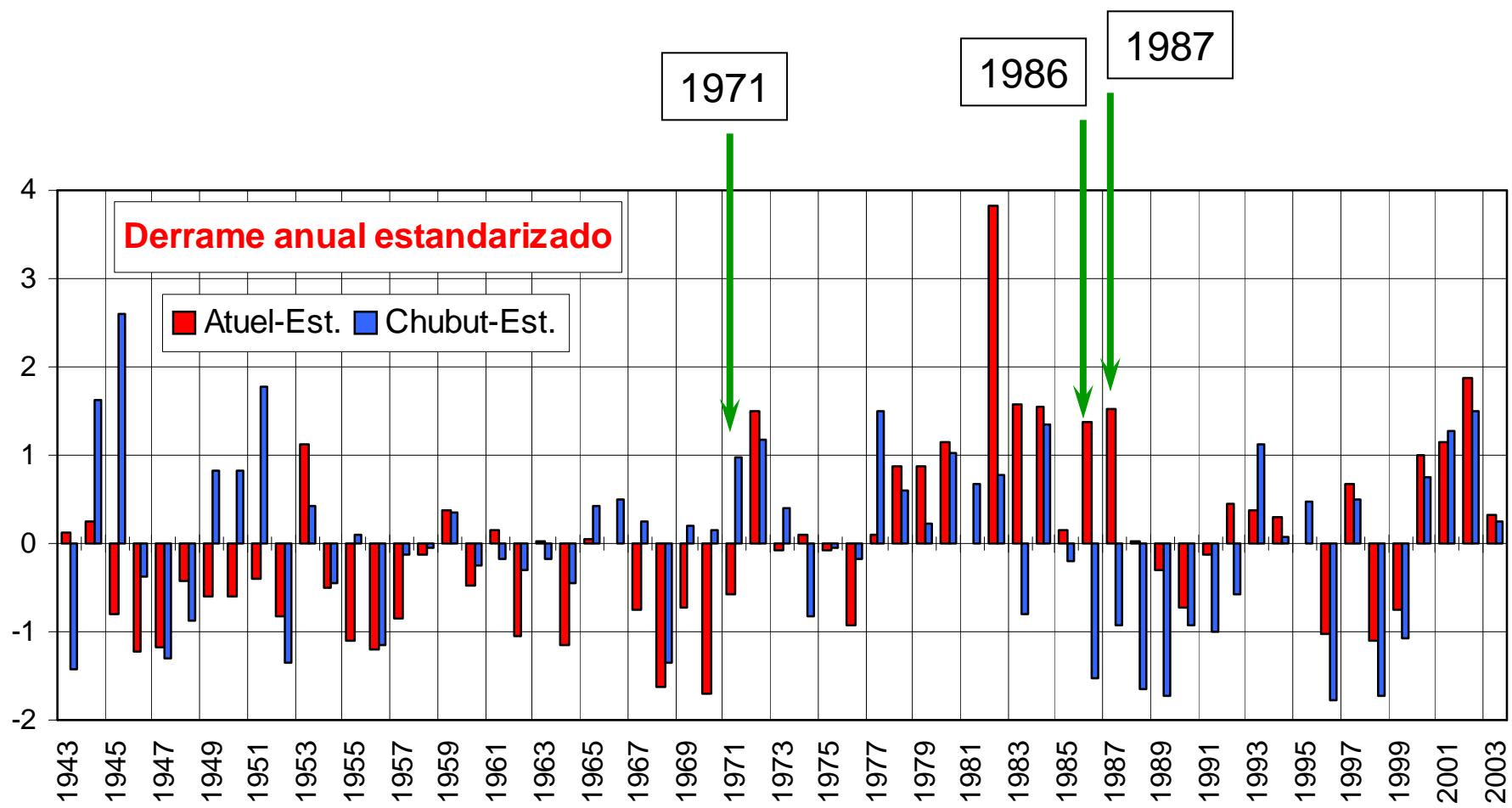
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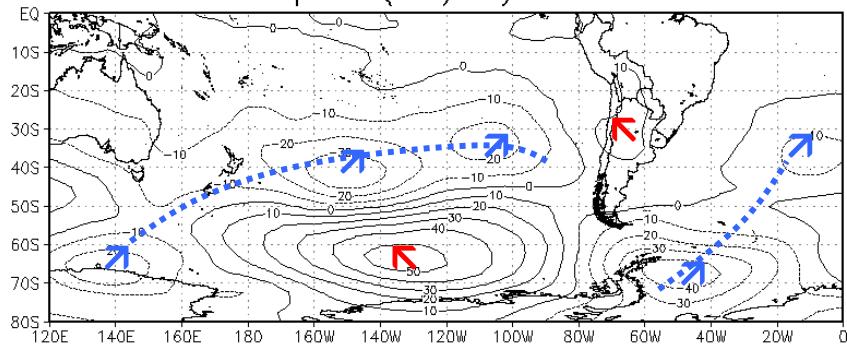




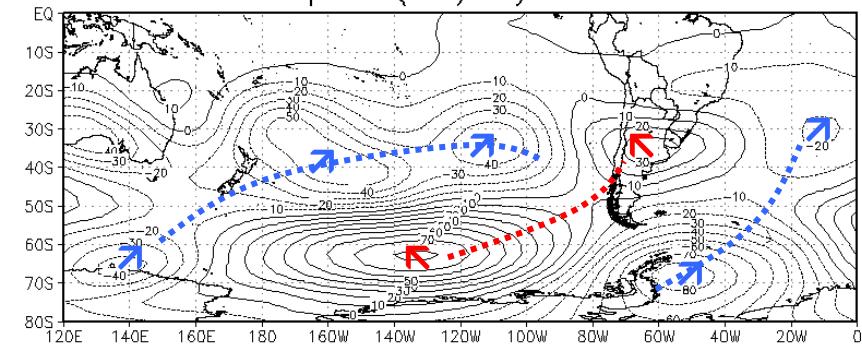




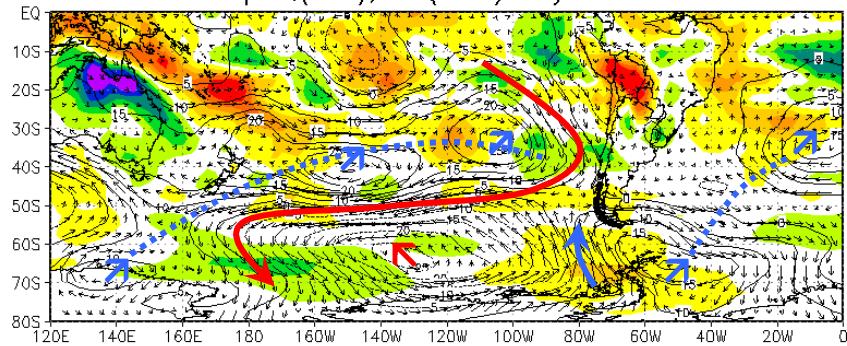
Comp: HGP(850) May–Oct 1986



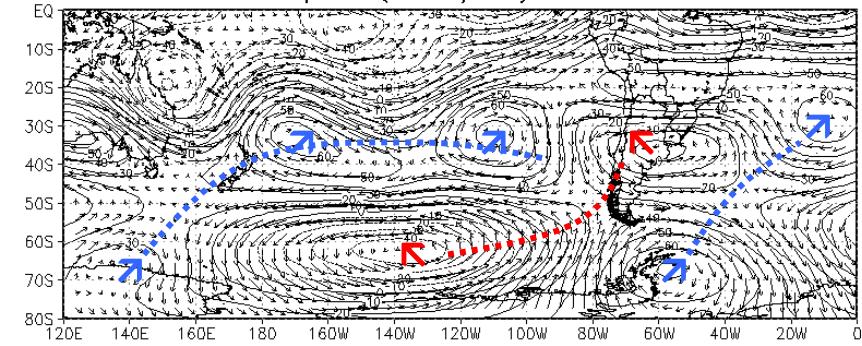
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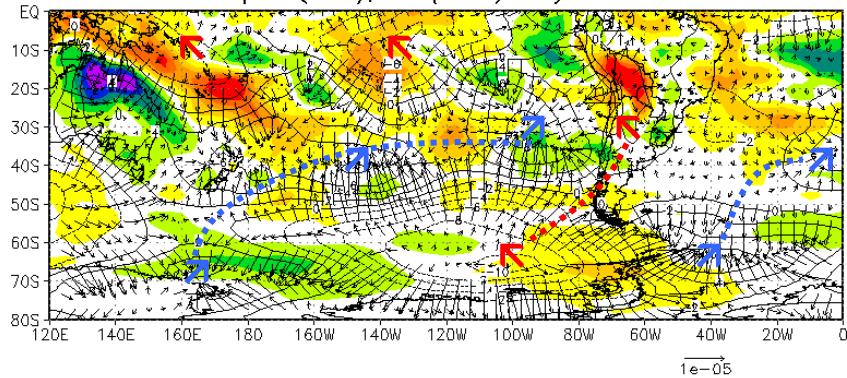
Comp: Q(850), Psi(.995) May–Oct 1986



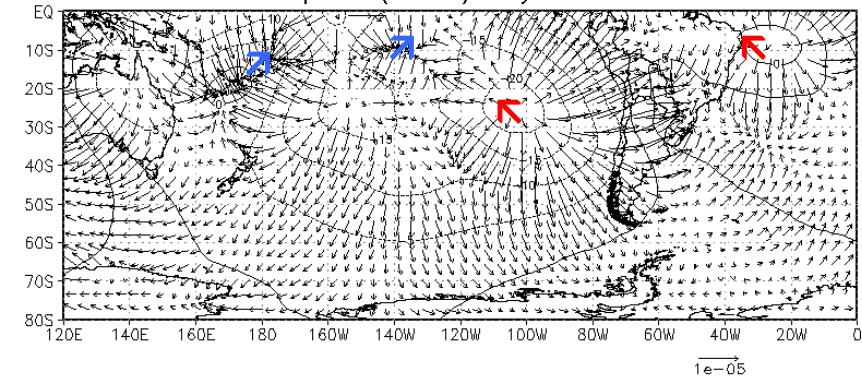
Comp: Psi(.2101) May–Oct 1986



Comp: Q(850), Chi(.995) May–Oct 1986

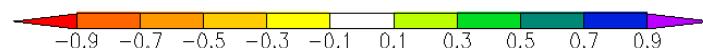
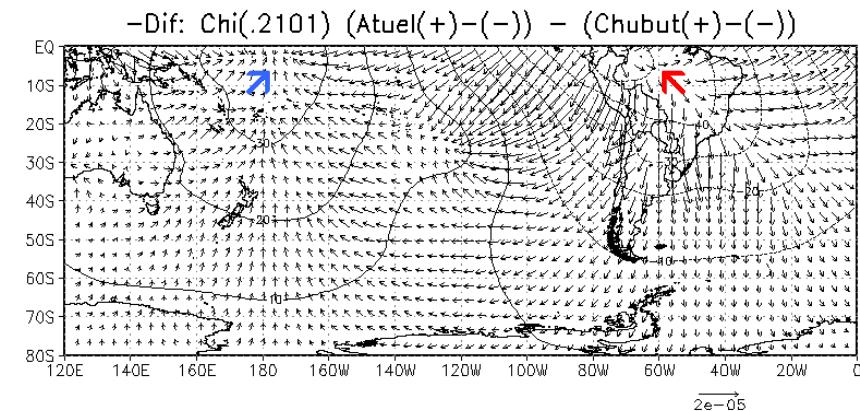
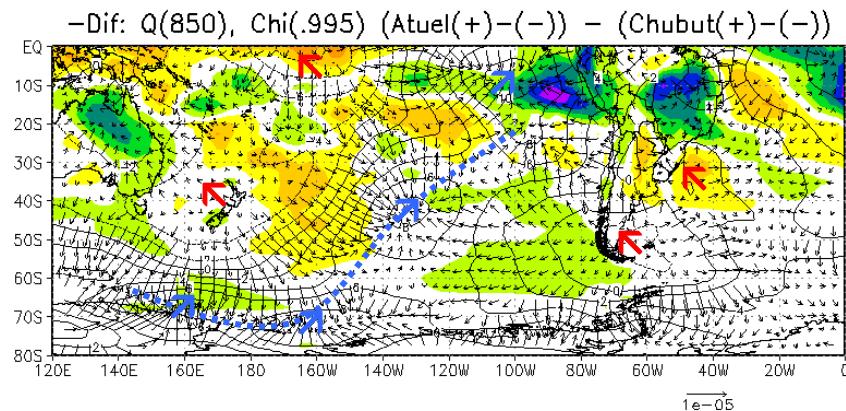
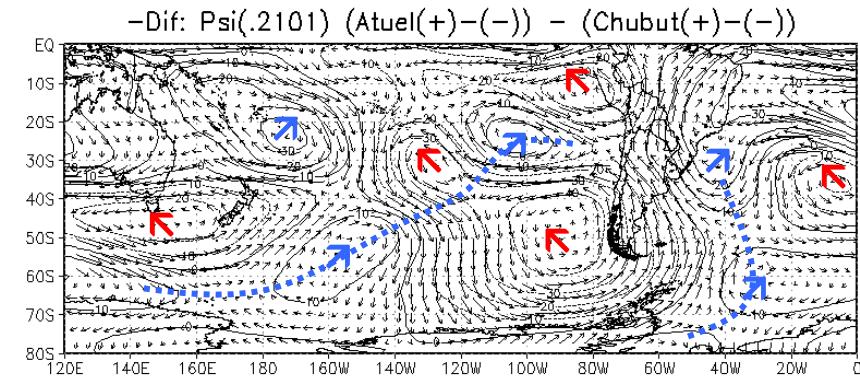
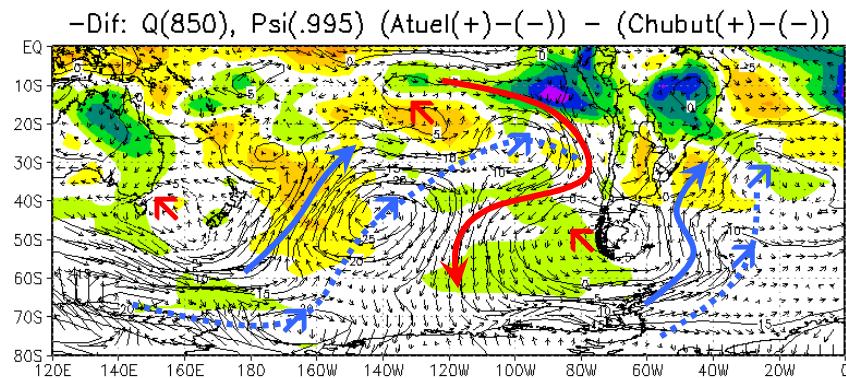
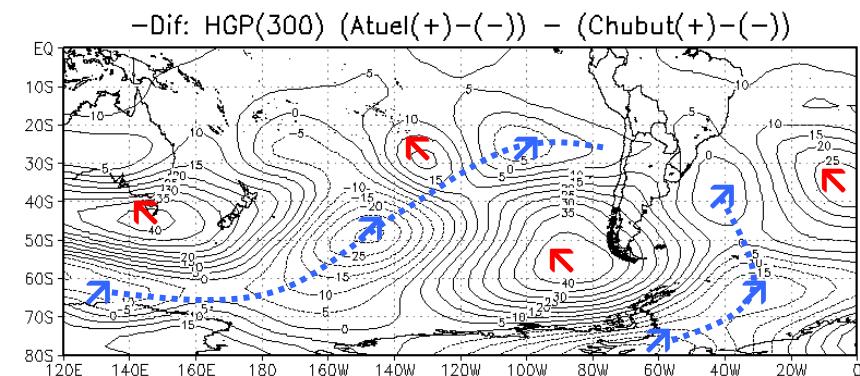
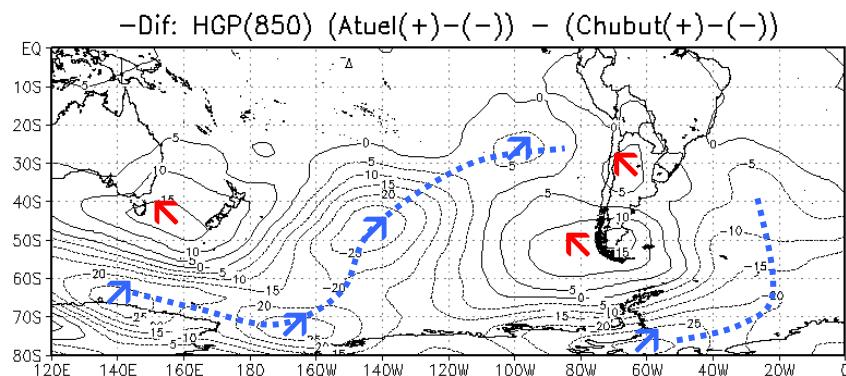


Comp: Chi(.2101) May–Oct 1986



-0.9 -0.7 -0.5 -0.3 -0.1 0.1 0.3 0.5 0.7 0.9

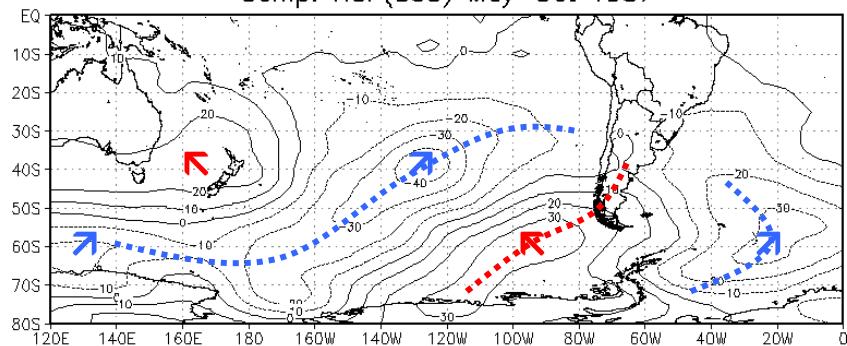
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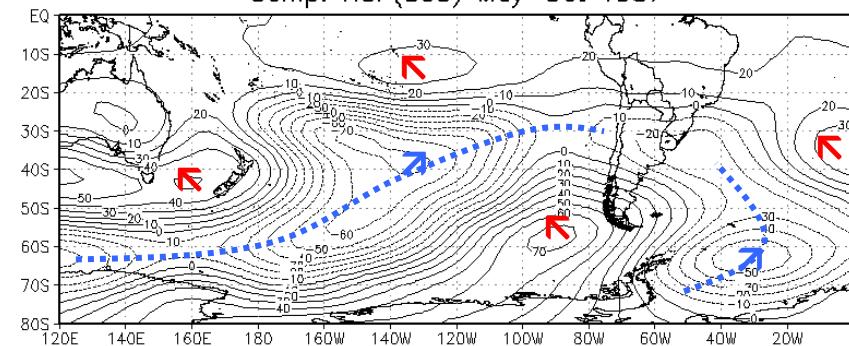
$1e-05$

$2e-05$

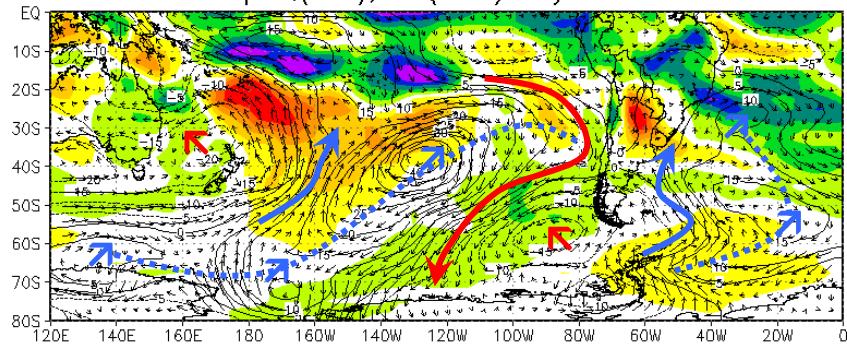
Comp: HGP(850) May–Oct 1987



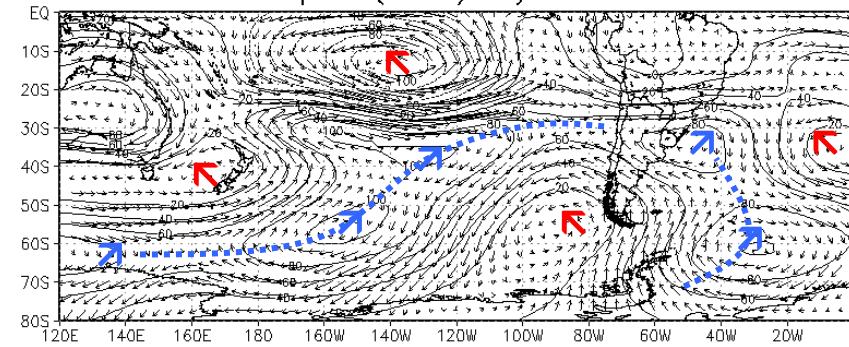
Comp: HGP(300) May–Oct 1987



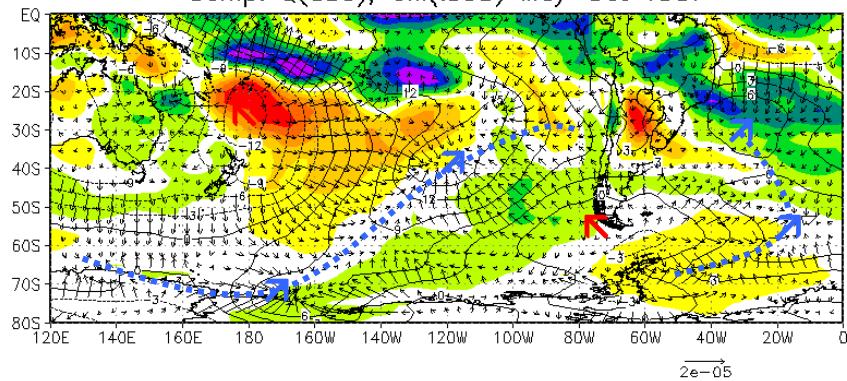
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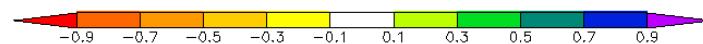
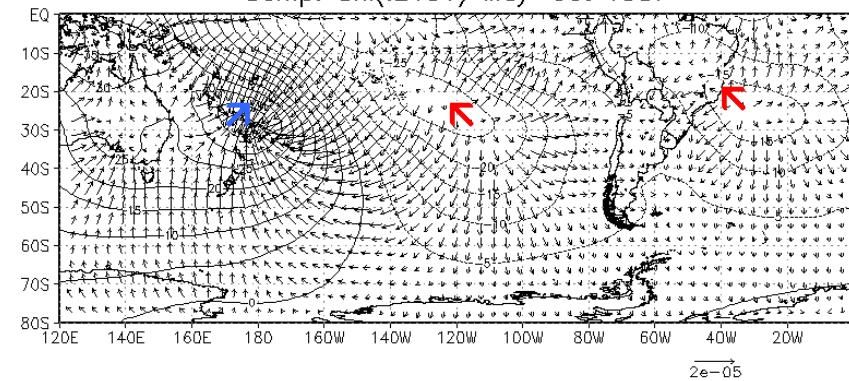
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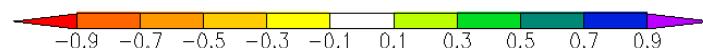
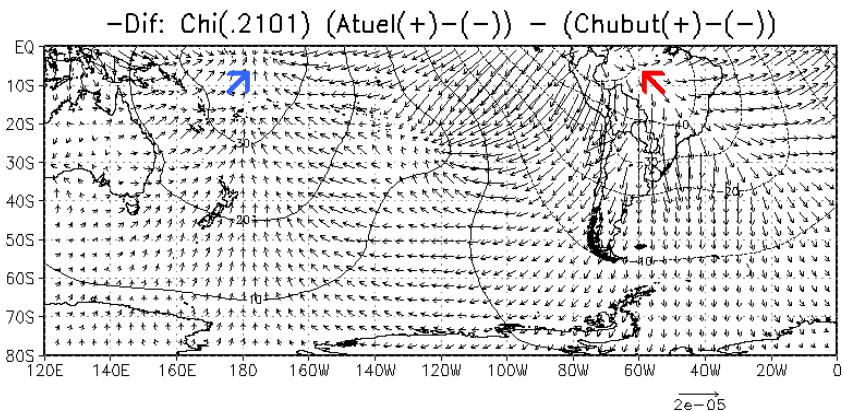
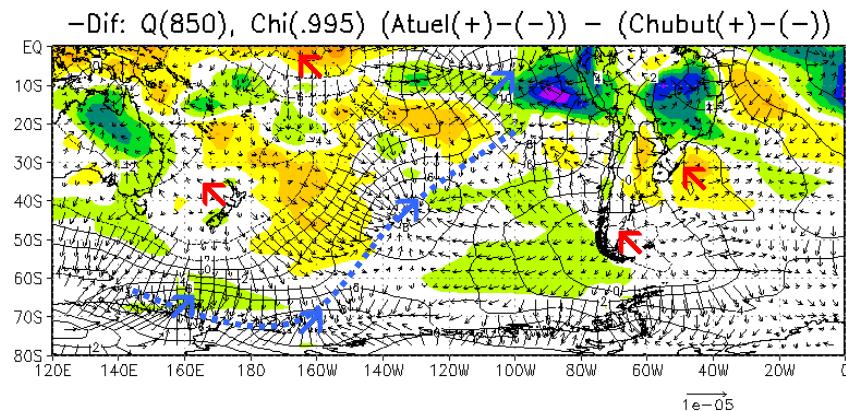
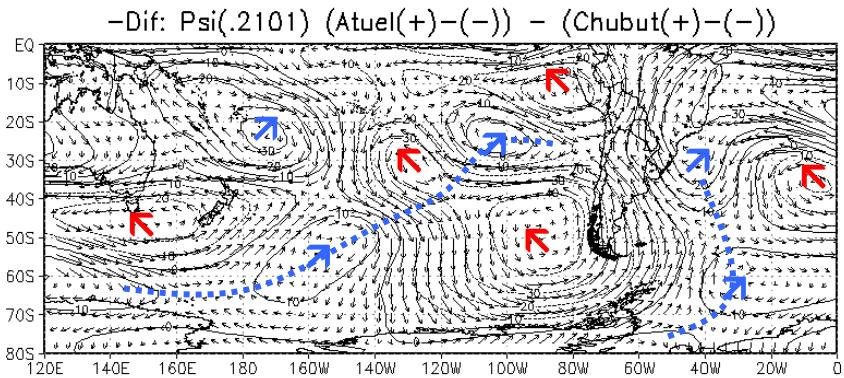
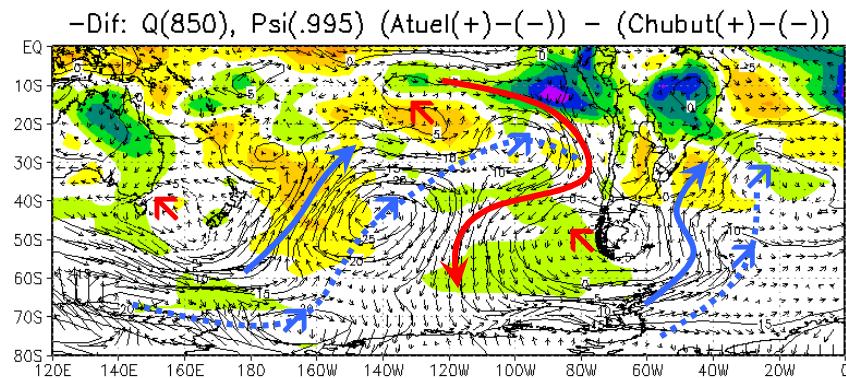
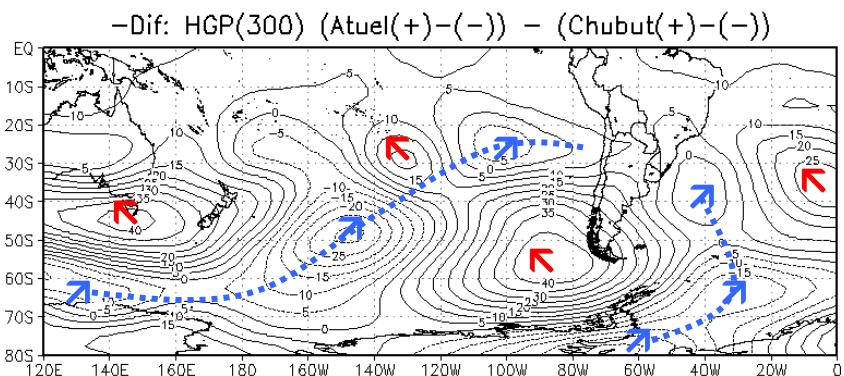
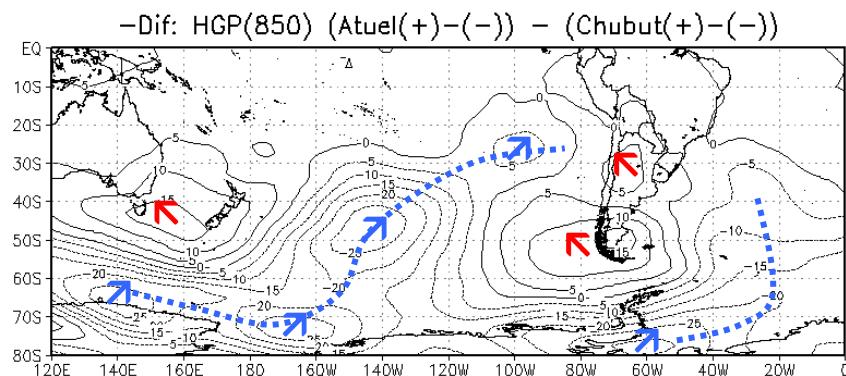
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Comp: Chi(.2101) May–Oct 1987



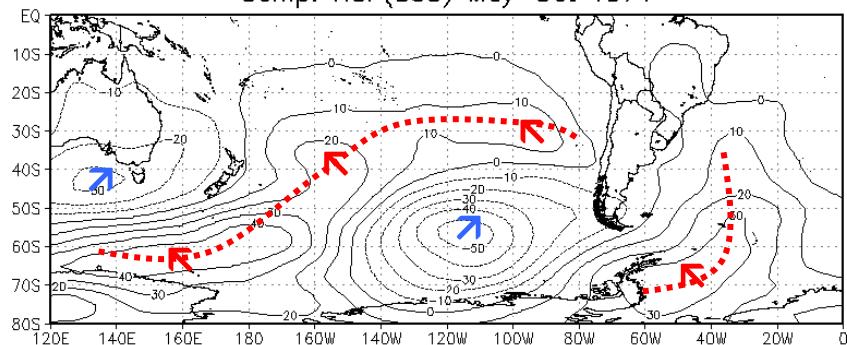
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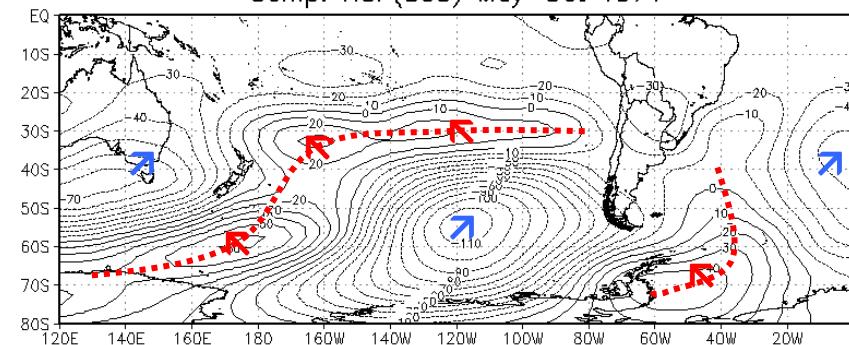
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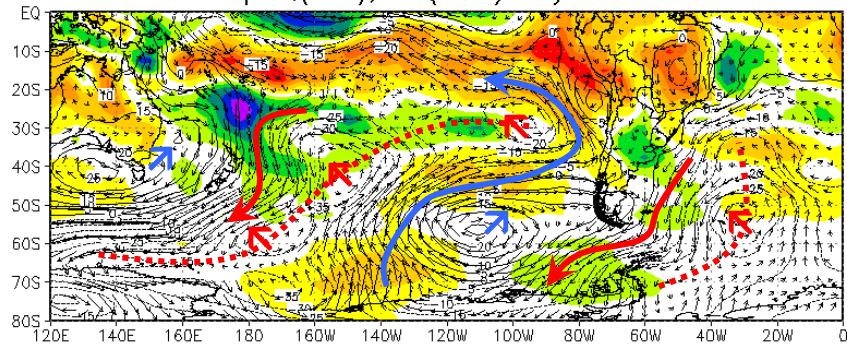
Comp: HGP(850) May–Oct 1971



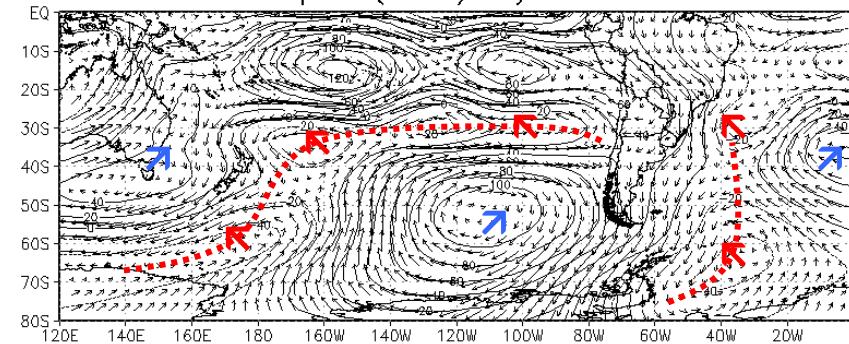
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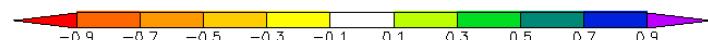
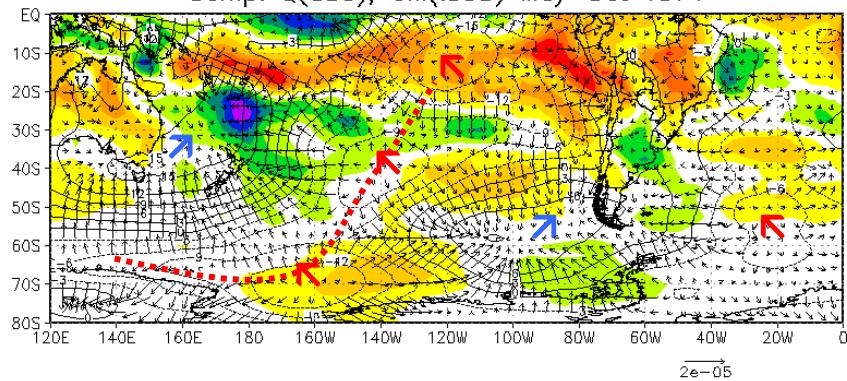
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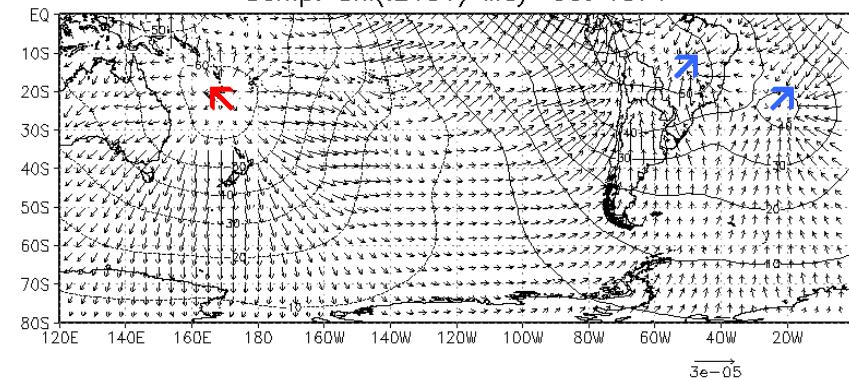
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Comp: Chi(.2101) May–Oct 1971

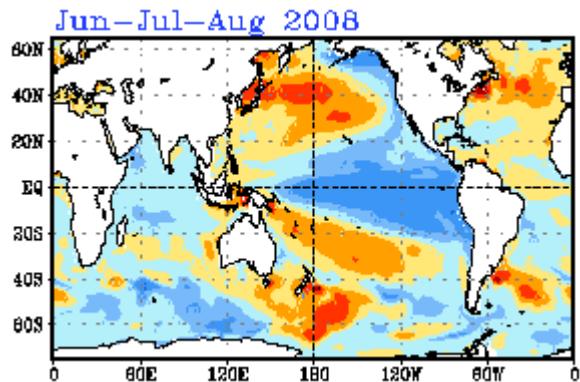


3×10^{-5}

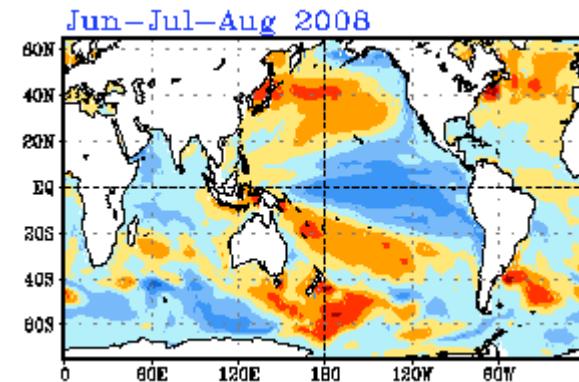


NWS/NCEP

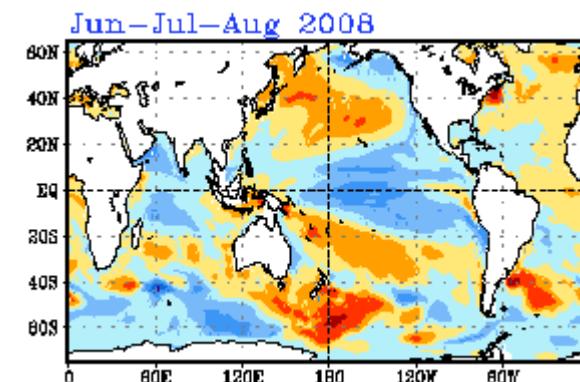
Temperatura Superficial del Mar



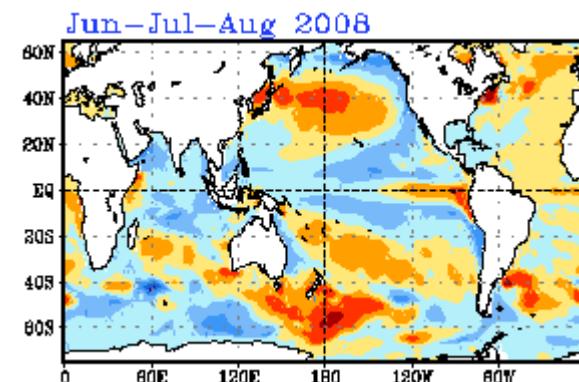
Enero 08



Febrero 08



Marzo 08



Abril 08

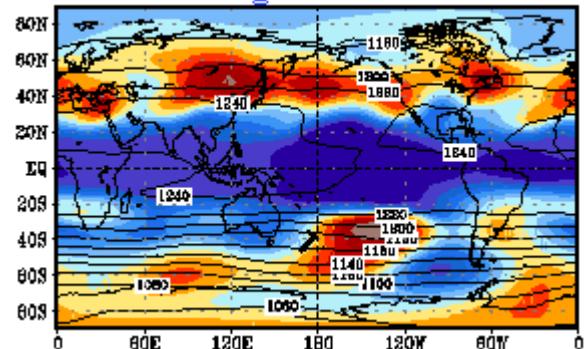




NWS/NCEP

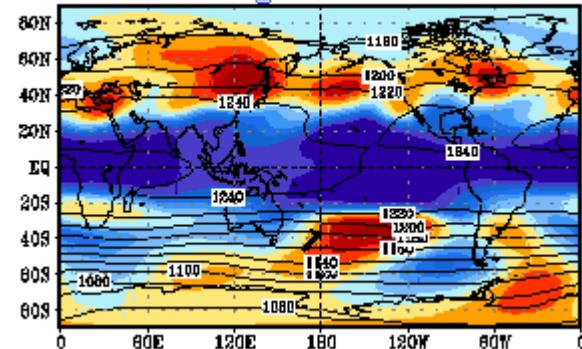
Altura Geopotencial 200 hPa

Jun-Jul-Aug 2008



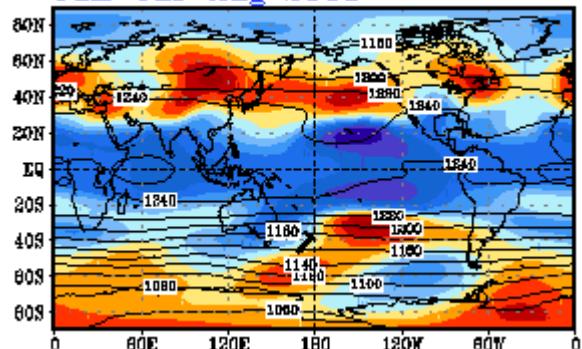
Enero 08

Jun-Jul-Aug 2008



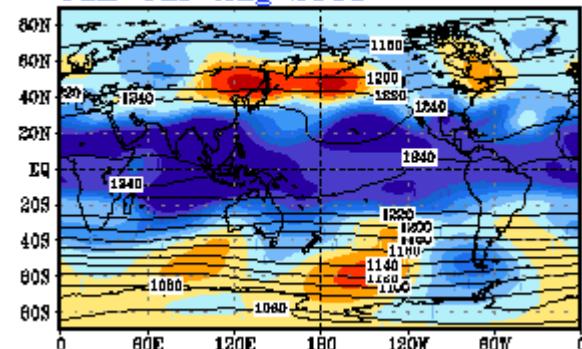
Febrero 08

Jun-Jul-Aug 2008

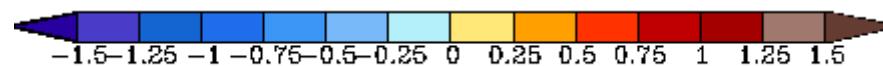


Marzo 08

Jun-Jul-Aug 2008



Abril 08

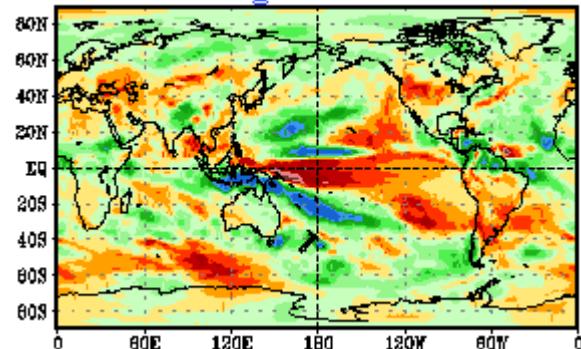




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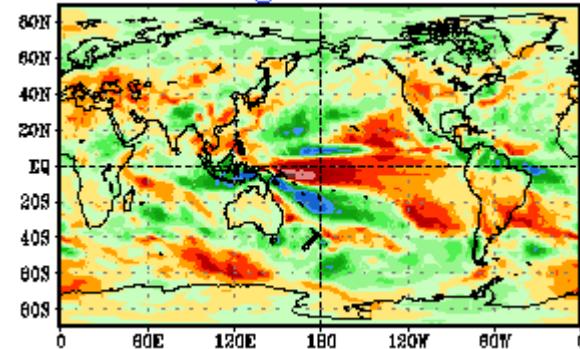
Precipitación

Jun-Jul-Aug 2008



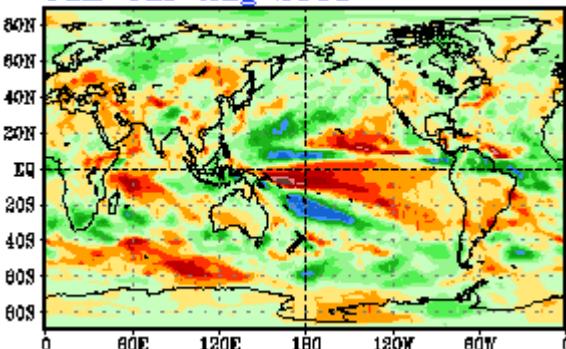
Enero 08

Jun-Jul-Aug 2008



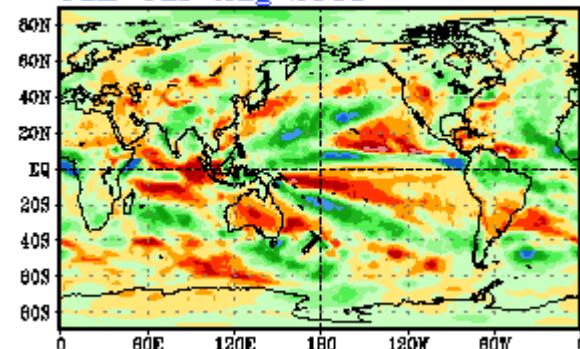
Febrero 08

Jun-Jul-Aug 2008

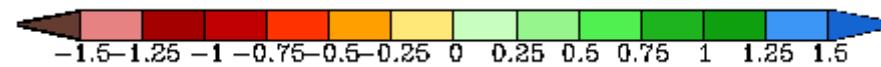


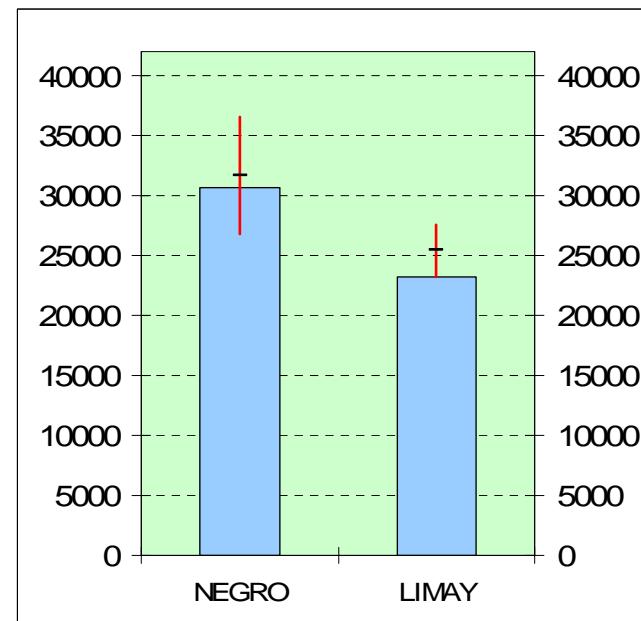
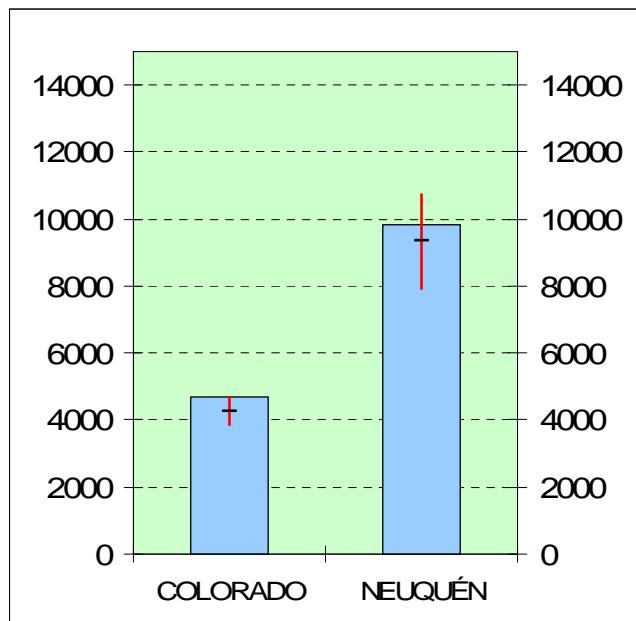
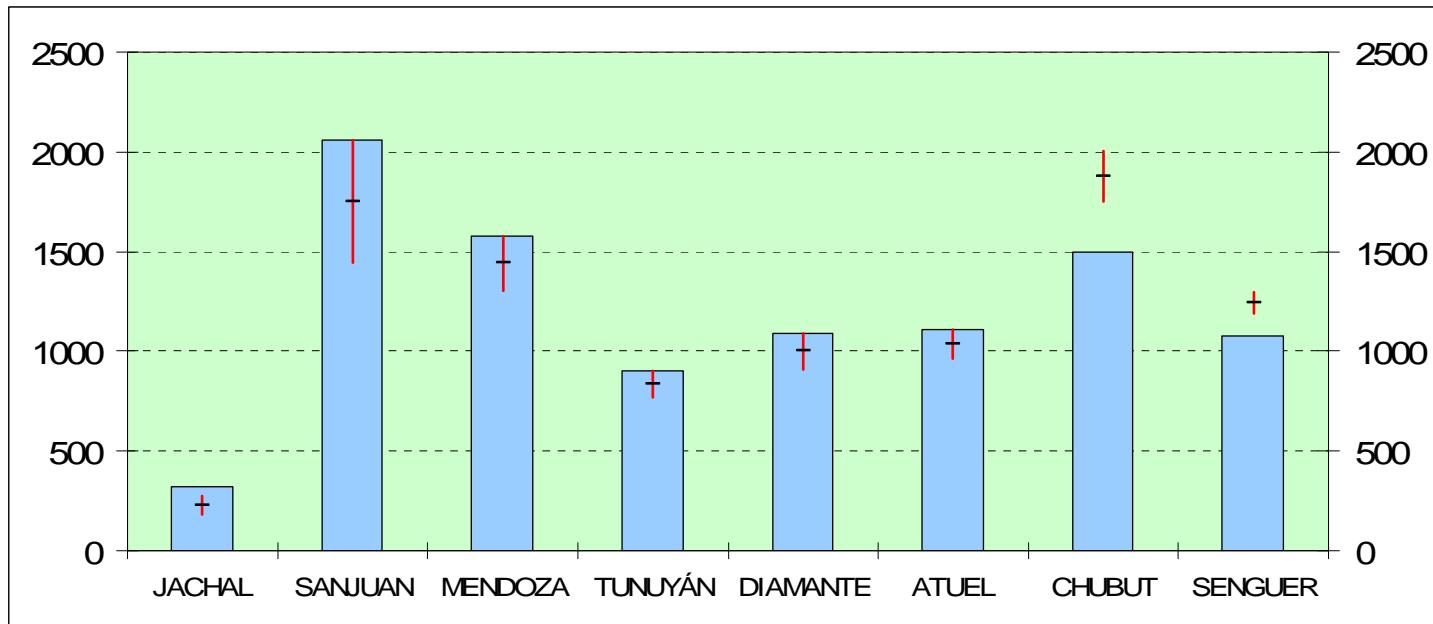
Marzo 08

Jun-Jul-Aug 2008



Abril 08





CONCLUSIONES

1. La variabilidad de baja frecuencia de los caudales de los ríos andinos se encuentra relacionada con las condiciones atmosféricas y oceánicas del Pacífico durante la estación fría.
2. Altos caudales se encuentran asociados a:
 - Debilitamiento del flujo de los O al O del Pasaje de Drake y de los anticiclones semipermanentes (debilitamiento del flujo básico).
 - Alteración de la circulación de Walker con una intensificación o corrimiento hacia el E de la ZCP (sobre el Pacífico S central).
 - Flujo de ondas de Rossby estacionarias (ORE) que atraviesa zonalmente el continente por latitudes subtropicales, con fuente sobre el Pacífico S central, y ayuda a mantener anomalías de baja presión frente a las costas de Chile (consecuente arrastre de humedad desde el Pacífico tropical y ascensos sobre la cordillera).
 - Flujo de ORE de menor intensidad que se propaga desde el Pacífico S occidental en dirección NO-SE hasta el Pasaje de Drake y luego SO-NE hacia el Atlántico, que ayuda a mantener las anomalías positivas de presión al SO del continente.

3. Bajos caudales se encuentran asociados a:
 - Intensificación del flujo de los O al O del Pasaje de Drake y profundización de los anticiclones semipermanentes (intensificación del flujo básico).
 - Intensificación de la circulación de Walker.
 - Flujo de ORE que se propaga desde el Pacífico S occidental en dirección NO-SE hasta el Pasaje de Drake y luego SO-NE hacia el Atlántico, que ayuda a mantener las anomalías negativas de presión al SO del continente.
4. La distinta ubicación de las fuentes de ORE tendría su origen en la intensificación o desplazamiento hacia el E de la ZCP.
5. La variabilidad de la ZCP podría tener su origen en el anómalo calentamiento diferencial de la superficie del mar (de allí la relación entre la variabilidad de los caudales y el ENSO).
6. La relación entre el índice Niño3+4 y los caudales se mantiene significativa con un máximo de antelación de hasta ~14 meses para el N de Cuyo y el S de la Patagonia, ~8 meses para el S de Cuyo y ~1 mes para el N de Patagonia, previos al período de máximos caudales.
7. Las diferencias entre los extremos de caudal de los ríos Atuel y Chubut, serían fundamentalmente consecuencia de la posición e intensidad de las anomalías de circulación al SO del continente.